

libRadtran

library for radiative transfer calculations
Edition 1.0 for libRadtran version 1.2
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This edition of the libRadtran documentation is consistent with version 1.2 of libRadtran.

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Preface

libRadtran is a library of radiative transfer routines and programs. It has evolved from the `uvspec` radiative transfer model. If you are not familiar with `uvspec`, please note the following: The central program of the libRadtran package is an executable called `uvspec` which can be found in the ‘`tools`’ directory. If you are interested in a user-friendly program for radiative transfer calculations, this is the main information you need to know. A description of `uvspec` is provided in the first part of this manual, and examples including various input files for different atmospheric conditions are provided in the examples directory.

The ‘`tools`’ directory also provides related utilities, like e.g. a mie program (`mie`), some utilities for the calculation of the position of the sun (`zenith`, `noon`, `sza2time`), a few tools for interpolation, convolution, and integration (`spline`, `conv`, `integrate`), and some other small tools.

The second part of the manual (which is only available in the developer manual) describes library routines which might be of interest if you want to write your own programs. The documentation is far from complete at present. Available are routines to read ASCII files, to do interpolations, integrations, convolutions, for Mie theory, to calculate the position of the sun, and some other stuff. More might be available somewhen in the future.

Please note that this document is by no means complete. It is under rapid development and major changes will take place.

Acknowledgements

Many people have already contributed to libRadtran’s development. In addition to Arve Kylling (arve.kylling@nilu.no) and Bernhard Mayer (bernhard.mayer@dlr.de), the following people have contributed to libRadtran or helped out in various other ways.

- The `disort` solver was developed by Knut Stamnes, Warren Wiscombe, S.C. Tsay, and K. Jayaweera
- Warren Wiscombe provided the Mie code `MIEV0`, and the routines to calculate the refractive indices of water and ice, `REFWAT` and `ICEWAT`.
- Seiji Kato (kato@aerosol.larc.nasa.gov) provided the correlated-k tables described in Kato et al. (1999).
- Tom Charlock (t.p.charlock@larc.nasa.gov), Quiang Fu (qfu@atm.dal.ca), and Fred Rose (f.g.rose@larc.nasa.gov) provided the most recent version of the Fu and Liou code.
- David Kratz (kratz@aquila.larc.nasa.gov) provided the routines for the simulation of the AVHRR channels described in Kratz (1995).
- Frank Evans (evans@nit.colorado.edu) provided the `polradtran` solver.
- Ola Engelsen provided data and support for different ozone absorption cross sections.
- Albano Gonzales (aglezf@ull.es) included the Yang et al. (2000), Key et al. (2002) ice crystal parameterization.
- Tables for the radiative properties of ice clouds for different particle “habits” were obtained from Jeff Key and Ping Yang, Yang et al. (2000), Key et al. (2002). In

addition, Ping Yang and Heli Wei kindly provided a comprehensive database of particle single scattering properties which we used to derive a consistent set of ice cloud optical properties for the spectral range 0.2 - 100 micron following the detailed description in Key et al. (2002).

- Paul Ricchiazzi (paul@icess.ucsb.edu) and colleagues allowed us to include the complete gas absorption parameterization of their model SBDART into uvspec.
- Luca Bugliaro (luca.bugliaro@dlr.de) wrote the analytical TZS solver (thermal, zero scattering).
- Sina Lohmann (sina.lohmann@dlr.de) reduced the “overhead time” for reading the Kato et al. tables dramatically which resulted in a speedup of a factor of 2 in a twostr solar irradiance calculation.
- Claudia Emde (claudia.emde@dlr.de) implemented detailed ice cloud properties provided by Bryan Baum (bryan.baum@ssec.wisc.edu).
- Yongxiang Hu (yongxiang.hu-1@nasa.gov) provided the delta-fit program used to calculate the Legendre coefficients for `ic_properties` `baum`.
- Many unnamed users helped to improve the code by identifying or fixing bugs in the code.

— The Detailed Node Listing —

1 A Brief Overview of libRadtran

This manual documents how to install and use libRadtran and corresponds to libRadtran version 1.2.

libRadtran is a collection of C and Fortran functions and subroutines useful for radiative transfer calculations in the Earth's atmosphere. In addition, programs in C, Fortran and Perl using the libRadtran functions and subroutines are included to allow the user to develop his/her own programs.

It is expected that the reader is familiar with radiative transfer terminology. In addition, a variety of techniques and parameterizations from various sources are used. For more information about the usefulness and applicability of these methods in certain contexts, the user is referred to the referenced literature.

1.1 Radiative transfer calculations

The central program of the libRadtran package is called **uvspec**. It was originally designed to calculate spectral irradiance in the ultraviolet and visible spectral ranges, hence the name. Over the years, **uvspec** has evolved to become a tool for many applications, including the simulation of instruments, the calculation of the radiation budget of the Earth, or the development of remote sensing techniques. **uvspec** is driven with a human-readable input file which allows the definition of the model input in a user-friendly way. Various commands are available to specify the properties of the atmosphere, including Rayleigh scattering, molecular absorption, aerosols, water and ice clouds, and the surface albedo. A selection of several radiative transfer solvers is available to simulate different aspects of the radiation field, including the **disort** solver and a pseudo-spherical version of this code, a fast twostream code, Frank Evans' **polRadtran**, and the three-dimensional **MYSTIC** code.

Radiative transfer calculations with **uvspec** are straightforward. The input to the radiative transfer solver is specified in the 'input_file'. Output is written to **stdout** and can easily be re-directed into an 'output_file':

```
uvspec < input_file > output_file
```

The syntax of input and output is described in [Section 2.1 \[uvspec\]](#), page 5. Examples of **uvspec** input and output files are found in the 'examples' directory.

1.2 Ozone retrieval from global irradiance measurements

Stamnes et al. (1991) described a method to derive the total ozone column from global irradiance measurements. The method is based on the comparison of measured irradiance ratios at two wavelengths in the UV part of spectrum with a synthetic chart of this ratio computed for a variety of ozone values. One of the wavelengths should be appreciably absorbed by ozone compared with the other. Typically choices are 305 and 340 nm. The method is reliable under cloudfree conditions, but increasingly overestimates the ozone column for optically thicker clouds, Mayer et al. (1998).

Within libRadtran the method is implemented with two pieces of software

Gen_o3_tab

Generates a look-up-table of ozone values as a function of the irradiance ratio for a given pair of wavelengths and solar zenith angle.

read_o3_tab

Takes as input the measured ratio and solar zenith angle, reads the look-up-table, and returns an ozone value.

The use of these tools is described in [Section 2.9 \[Geno3tab\]](#), page 59.

1.3 Cloud optical thickness from global irradiance measurements

Stamnes et al. (1991) described a method to derive a representative cloud optical depth from global irradiance measurements. The method compares the measured irradiance at a wavelength where ozone absorption is minimal to irradiances generated by a radiative transfer model as a function of cloud optical thickness. The method is very sensitive to the ground albedo and independent measurement of the albedo are needed when there is snow on the surface.

Within libRadtran the method is implemented with two pieces of software

Gen_wc_tab

Generates a look-up-table of water cloud optical depths for various solar zenith angles for a given wavelength.

read_o3_tab

Takes as input the measured irradiance and solar zenith angle, reads the look-up-table, and returns a cloud optical depth.

The use of these tools is described in [Section 2.10 \[Genwctab\]](#), page 60.

1.4 ... and much more

Apart from the above described programs, the ‘tools’ directory contains related utilities, like e.g. a mie program (**mie**), a tool to calculate wavelength-dependent cloud properties (**cldprp**), a simple shell script to add levels to existing profiles (**addlevel**), some utilities for the calculation of the position of the sun (**zenith**, **noon**, **sza2time**), a few tools for interpolation, convolution, and integration (**spline**, **conv**, **integrate**), and some other small utilities. Generally, these programs will give some information about their purpose and usage when called without arguments.

2 Some useful tools

2.1 uvspec

`uvspec` calculates the radiation field in the Earth's atmosphere for a given set of input parameters. It reads input from standard input, and outputs to standard output. It is normally invoked in the following way:

```
uvspec < input_file > output_file
```

The format of the input and output files are described below. Several realistic examples of input files are subsequently given.

`uvspec` may produce a wealth of diagnostic messages and warnings, depending on your use of `verbose` or `quiet`. Diagnostics, error messages, and warnings are written to `stderr` while the `uvspec` output is written to `stdout`. To make use of this extra information, you may want to write the standard `uvspec` output to one file and the diagnostic messages to another. To do so, try `(./uvspec < uvspec.inp > uvspec.out) >& verbose.txt`. The irradiances and radiances will be written to '`uvspec.out`' while all diagnostic messages go into '`verbose.txt`'. This method can also be used to collect `uvspec` error messages.

Warning: Please note the error checking on input variables is not complete at the moment. Hence, if you provide erroneous input, the outcome is unpredictable.

2.1.1 The uvspec input file

`uvspec` is controlled in a user-friendly way. The control options are named in a (hopefully) intuitive way. In the following, several examples are given, how to create an input file, how to define a cloudless sky atmosphere, how to add aerosols and clouds, etc. All following examples are taken from the `libRadtran` examples directory and are part of the `uvspec` self-check. For a complete listing and explanation of all input options, have a look at [section 2.1.3 \[Complete description of input parameters\]](#), page 21.

The `uvspec` input file consists of single line entries, each making up a complete input to the `uvspec` program. First on the line comes the parameter name, followed by one or more parameter values. The parameter name and the parameter values are separated by white space. Filenames are entered without any surrounding single or double quotes. Comments are introduced by a `#`. Blank lines are ignored. The order of the lines is not important, with one exception: if the same input option is used more than once, the second one will usually over-write the first one. Be aware that also option in another `included` input file will overwrite options specified before.

2.1.1.1 Cloudless, aerosol-free atmosphere

The simplest possible input file contains only a few lines:

```

                                # Location of atmospheric profile file.
atmosphere_file ../data/atmmod/afglus.dat

                                # Location of the extraterrestrial spectrum
solar_file ../data/solar_flux/atlas_plus_modtran

wavelength 310.0 310.0    # Wavelength range [nm]

```

The first three statements define the location of some data files: the uvspec data path (`data_files_path`) where all internal files are expected, the atmospheric profile (`atmosphere_file`), and the extraterrestrial spectrum (`solar_file`). The last line defines the desired wavelength range which is a monochromatic data point in this example. All other data which are not explicitly mentioned assume a default value which is "0" in most cases. Here, the solar zenith angle is 0, the surface albedo is 0, and the atmosphere does not contain clouds nor aerosols. Pressure, temperature, ozone concentration, etc. are read from `atmosphere_file`.

An example of a more complete input file for a clear sky atmosphere is

```

                                # Location of atmospheric profile file.
atmosphere_file ../data/atmmod/afglus.dat

                                # Location of the extraterrestrial spectrum
solar_file ../data/solar_flux/atlas_plus_modtran
ozone_column 300.            # Scale ozone column to 300.0 DU
day_of_year 170              # Correct for Earth-Sun distance
albedo 0.2                   # Surface albedo
sza 32.0                     # Solar zenith angle
rte_solver disort            # Radiative transfer equation solver
deltam on                    # delta-M scaling on
nstr 6                        # Number of streams
wavelength 299.0 341.0      # Wavelength range [nm]
slit_function_file ../examples/TRI_SLIT.DAT
                                # Location of slit function
spline 300 340 1             # Interpolate from first to last in step

quiet

```

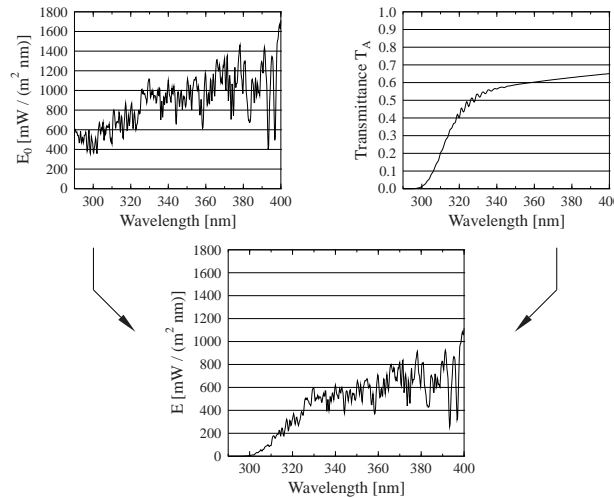
The atmosphere model, i.e. pressure, temperature, and ozone concentration profiles are read from `../data/atmmod/afglus.dat`. The extraterrestrial solar flux is read from the file `../data/solar_flux/atlas_plus_modtran`.

A wavelength dependent **surface albedo** may be specified using `albedo_file` instead of `albedo`. Non-Lambertian surface reflectance (BRDF) for vegetation and water may also be defined (please note that these require the use of `rte_solver disort2`). The BRDF of Vegetation is specified using `rpv_rho0`, `rpv_k`, and `rpv_theta`, following the definition of Rahman et al. (1993). Wavelength-dependent BRDF for vegetation can be defined with `rpv_file`. The BRDF of water surfaces is parameterized following Cox and Munk (1954a,

1954b) and Nakajima (1983). The respective parameters are the wind speed `cox_and_munk_u10`, the pigment concentration `cox_and_munk_pchl`, and the salinity `cox_and_munk_sal`. A complete description of these parameters is given in section [Section 2.1.3 \[Complete description of input parameters\]](#), page 21.

It is helpful to know some details about the **input/output wavelength resolution** in `uvspec` and how it can be influenced by the user. Basically there are three independent wavelength grids, the **input grid**, the **internal grid**, and the **output grid**. The essential thing to know is that the internal grid is chosen by `uvspec` itself in a reasonable way, if not explicitly defined in the input file with `transmittance_wl_file` or `molecular_tau_file`. The output grid is completely independent of the internal grid and is entirely defined by the `solar_file`. The wavelength grid of all other input data (e.g. albedo, optical properties of aerosols and clouds, etc) is also completely independent. These data are automatically interpolated to the resolution of the internal wavelength grid. Hence, only two constraints are set to the gridding of the input data: (1), the wavelength range has to cover all internal grid points; and (2), it should be chosen in a reasonable manner to allow reasonable interpolation (which essentially means, dense enough).

In the ultraviolet/visible, `uvspec` uses an internal grid with a step with of 0.5nm below 350nm and 1nm above 350nm. This is a conservative choice which fully resolves the broad ozone absorption bands and the slowly varying Rayleigh, aerosol, and cloud extinctions. The idea is outlined in the following figure which is taken from Mayer et al. (1997):



The transmittance (or reflectance) is calculated on a moderate resolution grid which reduces the number of calls to the `rte_solver` and hence the computational time. Then, the transmittance is interpolated to the wavelengths in the `solar_file` (which is usually defined with higher spectral resolution), multiplied with the extraterrestrial irradiance, and possibly post-processed. Hence, the wavelength in the output spectrum are those contained in the `solar_file` which has two important implications: (1) Only those wavelengths are output that are contained in the `solar_file`. If e.g. a monochromatic calculation is defined by setting '`wavelength 327.14 327.14`', there will only be output if the wavelength 327.14 is explicitly listed in `solar_file`; (2) this is also true at thermal wavelengths where the extraterrestrial irradiance is zero; hence, even for a calculation in the thermal range a `solar_`

`file` can be specified which defines the output grid in the first column and arbitrary values in the second column. Keeping these points in mind, `solar_file` is a convenient way to define an arbitrary output grid. `solar_file` may be omitted for thermal radiation calculations (`source thermal`) as well as for `transmittance` and `reflectivity` calculations. If omitted, the output grid equals the internal wavelength grid.

If required, a **user-defined internal grid** can be specified with `transmittance_wl_file` or `molecular_tau_file`. Note that this is a way to speed up the calculation considerably. E.g., for some applications the internal grid in the UV-A and visible can be set to 10nm which would reduce computational times by up to a factor of 10.

Things are completely different if one of the `correlated_k` parameterizations is selected (see below). In this case all flexibility is taken away from the user which is an inherent feature of the `k` distribution method. Internal grid as well as the extraterrestrial file are in this case defined by the choice of the parameterization itself.

2.1.1.2 Spectral resolution

`uvspec` offers four different ways of spectral calculations:

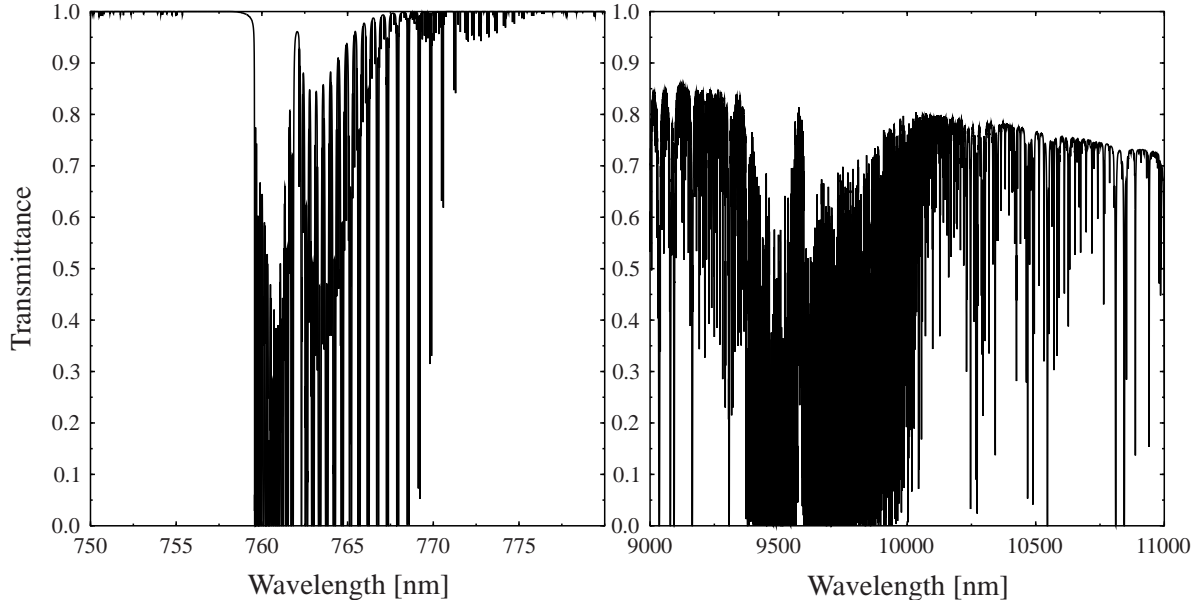
1. **Spectrally resolved calculation** in the UV and visible spectral ranges;
2. **Line-by-line calculation** with user-defined molecular absorption data;
3. **The correlated-k method**.
4. **Pseudo-spectral calculation** with exponential-sum-fit, from LOWTRAN; code adopted from SBDART (Ricchiazzi et al., 1998);

The choice of the method is determined by the problem and the decision is therefore entirely up to the user. The spectrally resolved calculation and the line-by-line calculation are more or less exact methods while the correlated-k distribution and the pseudo-spectral calculation are approximations that provide a compromise between speed and accuracy. In the following it is briefly described which method fits which purpose:

A **spectrally resolved calculation** is the most straightforward way, and will be the choice for all users interested in the ultraviolet and visible spectral ranges. In the UV/vis gas absorption generally occurs in broad bands with only slow spectral variation, the most important of these being the Hartley, Huggins, and Chappuis bands of ozone. Hence, a radiative transfer calculation every 1nm usually is sufficient to fully resolve any spectral variation using the method described in the last section. Absorption cross sections for various species are included, among them the most important O₃ and NO₂.

In the infrared, however, molecular absorption spectra are characterized by thousands of narrow absorption lines. There are two ways to treat these, either by highly resolved spectral calculations, so-called **line-by-line** calculations, or by a band parameterization. Concerning line-by-line, `uvspec` offers the possibility to define a spectrally resolved absorption cross section profile using `molecular_tau_file`. There is no option in libRadtran to generate such a `molecular_tau_file`, because (1) the HITRAN database which forms the basis for such calculations amounts to about 100 MByte which are updated continuously; and (2), there are sophisticated line-by-line programs available, like e.g. `genln2` (Edwards, 1992). Using `genln2` it is straightforward to create the input for `uvspec` line-by-line calculations. It is also planned to make line-by-line cross sections available for the six standard profiles that

come with libRadtran. The following figure shows an example of a line-by-line calculation of the atmospheric transmittance in two selected solar and thermal spectral ranges, the O2A-absorption band around 760nm and a region within the infrared window around 10 micron.

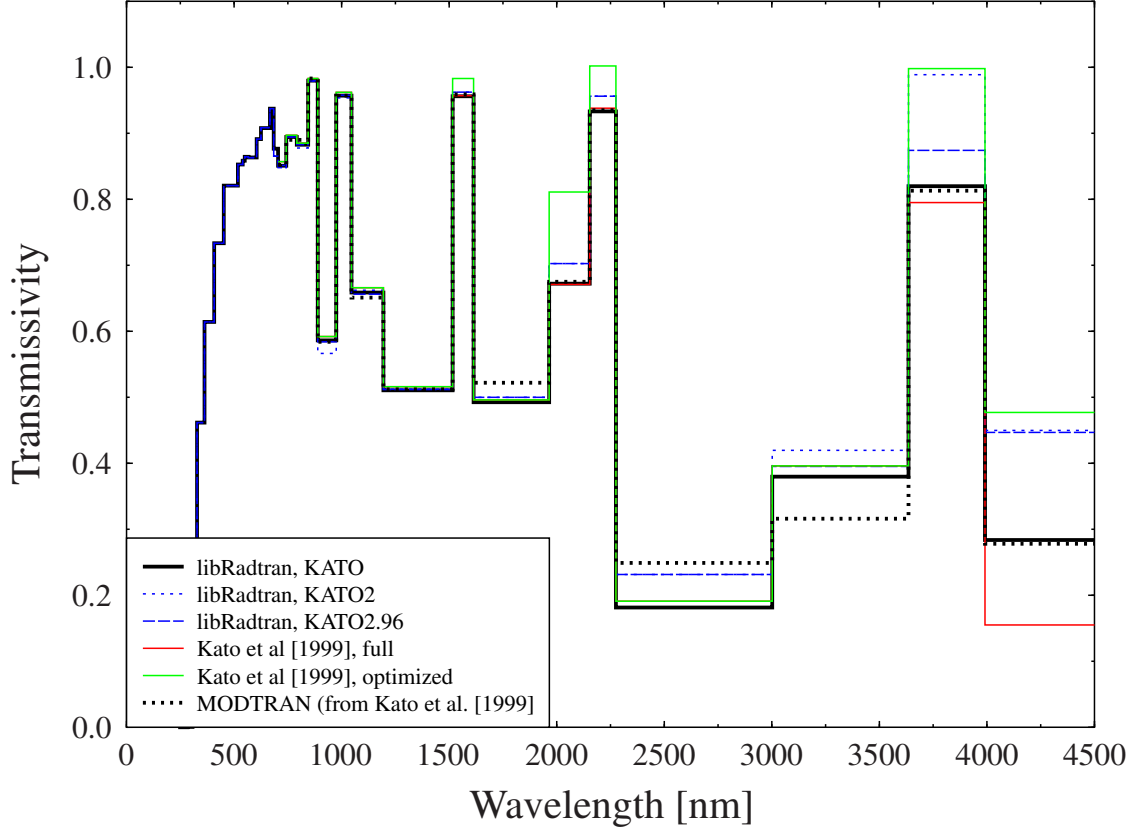


All spectral lines in the left figure are due to absorption by oxygen, while the ones in the left figure are due to ozone, water vapour, and CO₂. Line-by-line is obviously the exact way for radiation calculations. For most applications, however, line-by-line is far too slow. Here one needs a band parameterization, and the most accurate of these is the so-called **correlated-k approximation**. `uvspec` contains several correlated-k parameterizations which are invoked with `correlated_k`, in particular Kato et al. (1999), Fu and Liou (1992/93), Kratz (1995), as well as the possibility to specify a user-defined one. Kato et al. (1999) is a accurate parameterization for the solar spectral range. `uvspec` contains three different versions:

- | | |
|-----------------|---|
| Kato | The original tables provide by Seiji Kato which should correspond to the full version described in Kato et al. (1999); 575 subbands total, that is, 575 calls to the <code>rte_solver</code> |
| Kato2 | A new, optimized version of the tables, provided by Seiji Kato, 2003, with only 148 subbands (that is, calls to the <code>rte_solver</code>); the uncertainty is only slightly higher than Kato ; the absorption coefficients are based on HITRAN 2000. |
| Kato2.96 | Similar to Kato2 but based on HITRAN96. |

The following Figure shows a comparison between the three parameterization which are part of libRadtran and the data from Figure 3 by Kato et al. [1999]. It is immediately obvious that the uncertainty is high for all bands above 2.5 micron which is probably due to the treatment of band overlap. For this reasons, the results for the individual bands should not be trusted while the integrated shortwave radiation (the sum of all 32 bands) is calculated with high accuracy because (1) the bands above 2.5 micron contribute only

little to the integrated irradiance; and (2) errors are random and cancel each other to some degree.



For more information on these parameterizations please refer to the mentioned publications. Correlated-k is a powerful way to calculate spectrally integrated quantities, however, it takes away some flexibility. In particular, this means that the wavelength grid is no longer chosen by the user but by the parameterization, that is, by `uvspec`. The `uvspec` output is then no longer spectral quantities, e.g. $W / (m^2 nm)$, but integrated over the spectral bands, e.g. W / m^2 .

A simple but complete example for a correlated-k approximation of the solar spectrum:

```

# Conditions for the calculation of Figure 3 in
# Kato et al., JQSRT 62, 109-121, 1999.
# To compare the data, the direct irradiance calculated
# by uvspec has to be divided by cos(30 deg) because
# Kato et al. plot direct normal irradiance.

                                # Location of atmospheric profile file.
atmosphere_file ../examples/AFGLMS50.DAT
                                # Location of the extraterrestrial spectrum

albedo 0.2                      # Surface albedo
sza 30.0                       # Solar zenith angle
rte_solver twostr              # Radiative transfer equation solver

correlated_k KATO              # Correlated-k by Kato et al. [1999]

output sum                     # Calculate integrated solar irradiance

quiet

```

Here, the solar spectrum is split up into 32 bands according to Kato et al. (1999). In order to calculate integrated shortwave irradiance, simply sum the outputs, or even simpler, add `output sum` to the input file.

For **pseudo-spectral calculations** in the whole spectral range, we have adopted the molecular absorption parameterization from LOWTRAN/SBDART by Ricchiazzi et al. (1998). The respective section of this paper says:

SBDART relies on low-resolution band models developed for the LOWTRAN 7 atmospheric trans-mission code (Pierluissi and Peng, 1985). These models provide clear-sky atmospheric transmission from 0 to 50000 cm⁻¹ and include the effects of all radiatively active molecular species found in the earth's atmosphere. The models are derived from detailed line-by-line calculations that are degraded to 20 cm⁻¹ resolution for use in LOWTRAN. This translates to a wavelength resolution of about 5 nm in the visible and about 200 nm in the thermal infrared. These band models represent rather large wavelength bands, and the transmission functions do not necessarily follow Beers Law. This means that the fractional transmission through a slab of material depends not only on the slab thickness, but also on the amount of material penetrated before entering the slab. Since the radiative transfer equation solved by SBDART assumes Beers Law behavior, it is necessary to express the transmission as the sum of several exponential functions (Wiscombe and Evans, 1977). SBDART uses a three-term exponential fit, which was also obtained from LOWTRAN 7. Each term in the exponential fit implies a separate solution of the radiation transfer equation. Hence, the RT equation solver only needs to be invoked three times for each spectral increment. This is a great computational economy compared to a higher order fitting polynomial, but it may also be a source of significant error.

The LOWTRAN/SBDART gas parameterization is invoked with `correlated_k LOWTRAN`. The spectral resolution may be arbitrarily chosen by the user. If not explicitly

defined with `transmittance_wl_file`, an internal grid with a step width of 0.5nm below 350nm and 1nm above 350nm is chosen which is practically overkill for most applications in the infrared. An extraterrestrial spectrum covering the complete solar range is provided at two different resolutions, `data/solar_flux/kurudz_1.0nm.dat` and `data/solar_flux/kurudz_0.1nm.dat`. An example for the solar range is shown in `examples/UVSPEC_SBDART_SOLAR.INP`:

```
atmosphere_file ../data/atmmod/afglus.dat
solar_file ../data/solar_flux/kurudz_1.0nm.dat

albedo 0.2          # Surface albedo
sza 30.0           # Solar zenith angle

rte_solver twostr    # Radiative transfer equation solver
wavelength 250.0 2500.0 # Wavelength range

correlated_k SBDART  # select SBDART molecular absorption

aerosol_default
aerosol_visibility 20

quiet
```

while `examples/UVSPEC_SBDART_THERMAL.INP` shows how to do a thermal calculation:

```
# uvspec data files
data_files_path ../data/
atmosphere_file ../examples/AFGLUS.70KM
solar_file ../examples/UVSPEC_SBDART_THERMAL.TRANS

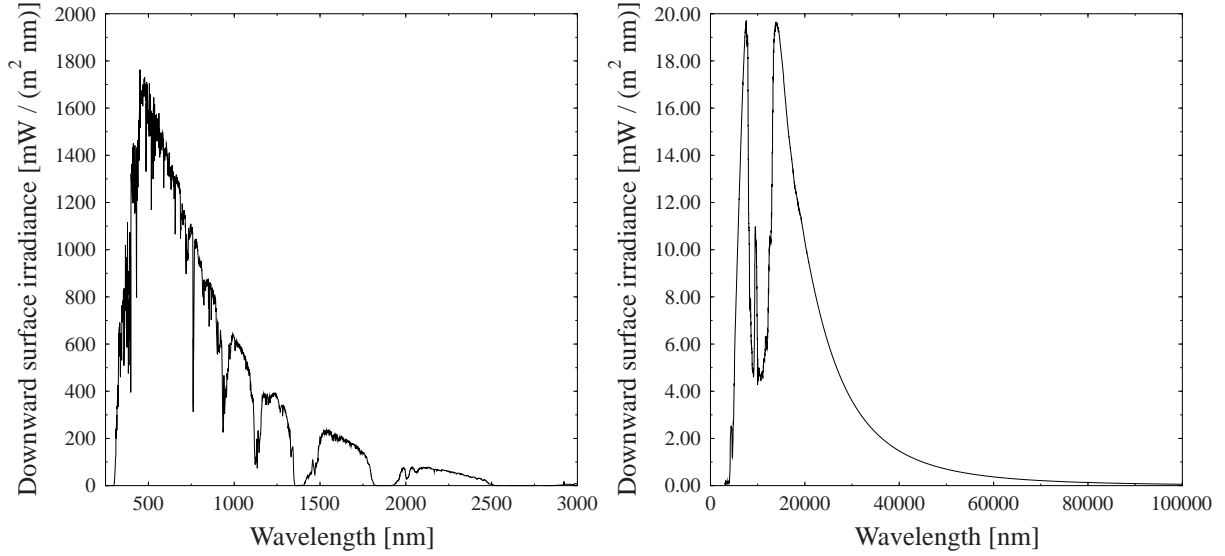
source thermal

rte_solver twostr    # Radiative transfer equation solver
transmittance_wl_file ../examples/UVSPEC_SBDART_THERMAL.TRANS

correlated_k SBDART  # select SBDART molecular absorption

output per_nm

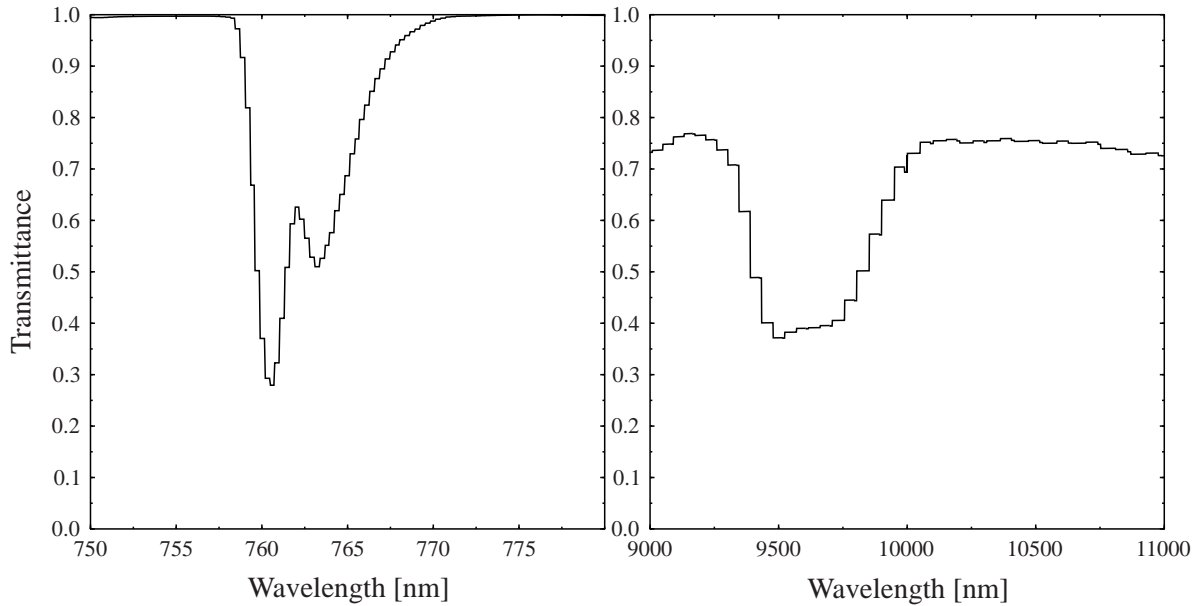
quiet
```



The figure shows the results of the solar and thermal calculations. The water vapour absorption bands in the solar range are clearly visible, as is the absorption window around 10 micron in the thermal. Please note the following points:

- Thermal radiation is per default output in W/(m² cm⁻¹), if the bandwidth is equal to 1 cm⁻¹ (default for `correlated_k` LOWTRAN calculations). Otherwise the output is the integrated flux over the wavenumber interval specified by `thermal_bandwidth`, `thermal_bands_file`, or by the `correlated_k` option (Kato, Kato2, Kato2.96, Fu, AVHRR_KRATZ, or Generic). To convert e.g. to W/(m² nm) use `output_per_nm` or multiply with k/λ where k is the wavenumber [cm⁻¹] and λ is the wavelength [nm]. To calculate band-integrated thermal quantities please consider `thermal_bands_file`.
- Even though no extraterrestrial irradiance is required, a `solar_file` may be specified for the thermal case. The reason is that, as explained initially, the `solar_file` defines the output grid. The second column in `solar_file` can be chosen arbitrarily in this case because it is ignored.
- For the choice of the wavelength grid for the calculation (`transmittance_wl_grid`) please consider that the resolution of the absorption parameterization is 5cm⁻¹ which converts to 0.3nm at 750nm and to 50nm at 10 micron. Choosing higher resolutions for the internal wavelength grid (`transmittance_wl_file`) is usually a waste of computational time.
- Please also make sure to choose a fine enough spectral resolution in order to capture all absorption features.

The following figure shows two selected wavelength intervals of the solar and thermal range, to demonstrate the spectral resolution of the LOWTRAN/SBDART absorption parameterization.



The resolution is about 5cm⁻¹ which translates to about 0.3nm in the left figure (oxygen A-band) and 50nm in the right figure (ozone absorption band in the atmospheric window). Compare this figure to the above line-by-line example to get an impression about the differences between both methods.

2.1.1.3 Aerosol

All options to set up and modify aerosol properties start with `aerosol_`. Aerosols may be specified in a hierarchical way. The most simple way to define an aerosol is by the command `aerosol_default` which will set up the aerosol model by E.P. Shettle, "Models of aerosols, clouds and precipitation for atmospheric propagation studies", in "Atmospheric propagation in the uv, visible, ir and mm-region and related system aspects", AGARD Conference Proceedings (454), 1989. The default properties are a rural type aerosol in the boundary layer, background aerosol above 2km, spring-summer conditions and a visibility of 50km. These settings may be modified with `aerosol_haze`, `aerosol_vulcan`, `aerosol_season`, and `aerosol_visibility`. More information can be introduced step by step, overwriting the default parameters. `aerosol_tau_file`, `aerosol_ssa_file`, and `aerosol_gg_file`, can be used to define the profiles of optical thickness, single scattering albedo, and asymmetry parameter. The integrated optical thickness can be set to a constant value using `aerosol_set_tau` or scaled with `aerosol_scale_tau`. The single scattering albedo may be scaled by `aerosol_scale_ssa` or set to a constant value by `aerosol_set_ssa`. The aerosol asymmetry factor may be set by `aerosol_set_gg`. The wavelength dependence of the aerosol optical depth is specified using the `aerosol_angstrom` parameter. `aerosol_moments_file` allows specification of the scattering phase function. If microphysical properties are available these may be introduced by defining the complex index of refraction `aerosol_refrac_index` or `aerosol_refrac_file` and the size distribution `aerosol_sizedist_file`. Finally, one may define the aerosol optical properties of each layer explicitly using `aerosol_files`.

The following list is an overview of the aerosol description parameters. The entries are arranged in a way that a parameter 'overwrites' all values higher up in the list.

aerosol_default

Generate default aerosol according to Shettle (1989)

aerosol_vulcan, aerosol_haze, aerosol_season, aerosol_visibility

Set Shettle (1989) aerosol properties (aerosol type, visibility)

aerosol_files

Specify optical properties of each layer explicitly, that is, extinction coefficient, single scattering albedo, and the moments of the phase function (everything as a function of wavelength).

aerosol_tau_file, aerosol_ssa_file, aerosol_gg_file

Overwrite profiles of optical thickness, single scattering albedo, and asymmetry parameter

aerosol_moments_file

Specify a phase function to be used instead of the Henyey-Greenstein phase function

aerosol_refrac_index, aerosol_refrac_file, aerosol_sizedist_file

Calculate optical properties from size distribution and index of refraction using Mie theory. Here is an exception from the rule that ALL values defined above are overwritten because the optical thickness profile is re-scaled so that the optical thickness at the first internal wavelength is unchanged. It is done that way to give the user an easy means of specifying the optical thickness at a given wavelength.

aerosol_set_gg, aerosol_set_ssa, aerosol_scale_ssa, aerosol_set_tau, aerosol_scale_tau

Overwrite profiles of asymmetry parameter and single scattering albedo

aerosol_angstrom

Overwrite the integrated optical thickness (profiles are not changed).

An example for a uvspec aerosol description is

```
include ../examples/UVSPEC_CLEAR.INP

aerosol_vulcan 1      # Aerosol type above 2km
aerosol_haze 6        # Aerosol type below 2km
aerosol_season 1      # Summer season
aerosol_visibility 20.0 # Visibility
aerosol_angstrom 1.1 0.2 # Scale aerosol optical depth
                        # using Angstrom alpha and beta
                        # coefficients
aerosol_scale_ssa 0.85 # Scale the single scattering albedo
                        # for all wavelengths
aerosol_set_gg 0.70    # Set the asymmetry factor
aerosol_tau_file ../examples/AERO_TAU.DAT
                        # File with aerosol optical depth profile
```

By combining this with the clear sky example given above a complete uvspec input file including aerosol is constructed.

2.1.1.4 Water clouds

All options to set up and modify water cloud properties start with `wc_`.

The easiest way to define a water cloud is to specify a `wc_file` which defines the liquid water content and effective droplet radius at each model level. By combining the following lines with the clear sky example given above a complete uvspec input file including water clouds is constructed.

```
include ../examples/UVSPEC_CLEAR.INP

wc_file ../examples/WC.DAT # Location of water cloud file
wc_set_tau 15.             # Set total water cloud optical depth
```

A typical example for a `wc_file` looks like:

#	z	LWC	R_eff
#	(km)	(g/m ³)	(um)
	5.000	0	0
	4.000	0.2	12.0
	3.000	0.1	10.0
	2.000	0.1	8.0

The three columns are the level altitude [km], the liquid water content [g/m³], and the effective droplet radius [micron]. Per default, these quantities are defined at the given model levels, that is, the value 0.2 g/m³ refers to altitude 4.0km, as e.g. in a radiosonde profile. The properties of each layer are calculated as average over the adjacent levels. E.g. the single scattering properties for the model layer between 3 and 4km are obtained by averaging over the two levels 3km and 4km. To allow easy definition of sharp cloud boundaries, clouds are only formed if both liquid water contents above and below the respective layer are larger than 0. Hence, in the above example, the layers between 2 and 3 as well as between 3 and 4km are cloudy while those between 1 and 2km and between 4 and 5km are not.

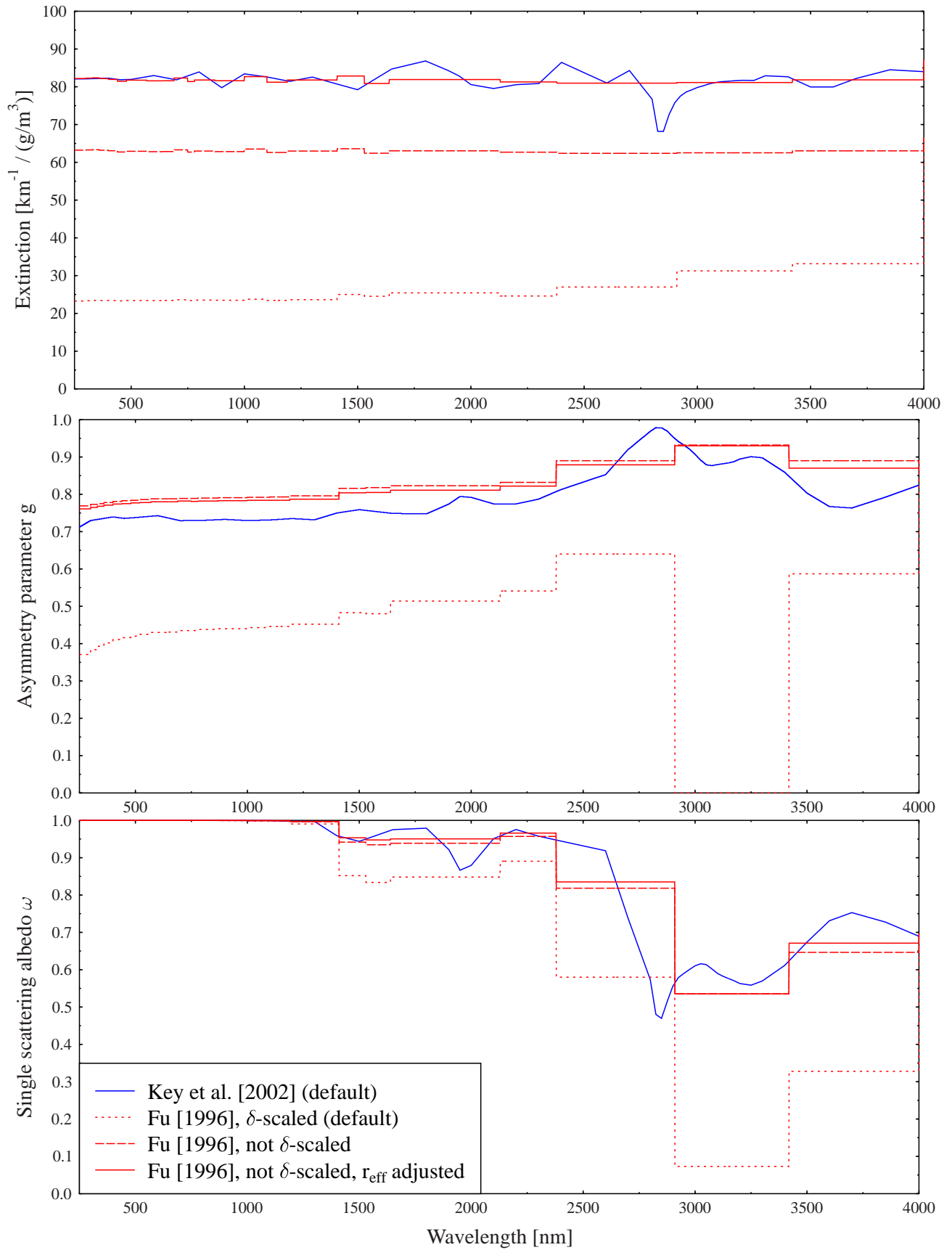
`wc_layer` provides a useful alternative if one prefers to define layer quantities. If specified, the properties are interpreted as layer properties, and in the above example, the cloud would extend from 2 to 5km, with e.g. a LWC of 0.2 g/m³ for the layer between 4 and 5km. To make sure that the clouds really look as you want them to look, it is recommended to use the `verbose` option. This option shows not only where the cloud is actually placed, it rather tells the user exactly how LWC and effective radius are translated into optical properties, depending on the choice of parameterisation. Please also note that the definition of the empty top level at 5km is important to tell uvspec where the cloud ends. If omitted, the cloud would extend all the way to the top of the atmosphere.

There are different ways to convert the microphysical properties to optical properties. Either a parameterization is used, like the one by Hu and Stamnes (1993) (which is the default), or by Mie calculations. The latter are very time-consuming, hence we decided not to include these online into `uvspec` but rather have an option to read in pre-calculated Mie tables. The option `wc_properties` controls the method: `hu` selects the Hu and Stamnes (1993) parameterization, `mie` selects pre-calculated Mie tables which are available at <http://www.libradtran.org>. If `wc_properties mie` is selected, the model expects one Mie cloud property file for each internal wavelength which is useful for the fixed wavelength grids used by the correlated-k parameterisations `correlated_k kato`, `correlated_k fu`, etc. For a spectral calculation with free wavelength grid, there is also the possibility to use a pre-defined set of Mie tables (available at the web site) and to define `wc_properties_interpolate` to automatically interpolate the Mie properties to the internal wavelength grid. Although this is an extremely useful option, please use it careful because it might consume enormous amounts of memory. Finally, there is the option to define an arbitrary file which is in the format as generated by Frank Evans' `cloudprp` which comes with SHDOM, see <http://nit.colorado.edu/~evans/shdom.html>.

As for the aerosol, there are several options to modify the optical properties of the clouds. And of course there is also the option of defining all cloud properties explicitly using `wc_files`.

2.1.1.5 Ice clouds

Ice clouds are generated in a similar way to water clouds. All options to set up and modify ice cloud properties start with `ic_`. The main difference between water and ice clouds is that the latter usually consist of non-spherical particles. Hence, the conversion from microphysical to optical properties is much less defined, and several parameterizations are available. Please note in addition that there are different definitions of the effective radius and e.g. the parameterizations by Key et al. (2002) and by Fu (1996) actually use different definitions (see explanation of `ic_properties`). Finally, the sharp forward peak which is typical for ice particles is also treated differently: E.g. Fu (1996) provides delta-scaled optical properties while Key et al. (2002) uses unscaled parameters (see explanation of `ic_fu_tau`). The following figure illustrates the implications. Plotted are extinction coefficient, asymmetry parameter, and single scattering albedo for ice clouds as a function of wavelength. If treated consistently, both Key et al. (2002) and Fu (1996) provide nearly identical results (blue and red solid line). However, Fu (1996) uses a different definition of effective radius and recommends to use the delta-scaled properties (red dotted line). The larger part of the difference compared to the solid line is due to delta-scaling which can be seen by comparing the red dotted (delta-scaled) and red-dashed (not delta-scaled) lines. Lately, a new set of optical properties was derived, combining the Key et al. (2002) data below 3.4 micron with new calculations above 3.4 micron, to be invoked with `ic_properties yang`.



2.1.1.6 Calculation of radiances

To calculate radiances the following lines will do the job when combined with the clear sky example above

```
include ../examples/UVSPEC_AEROSOL.INP # Include's may be nested.

rte_solver disort2 # This override what is specified in above file
                  # and files included in that file etc.

phi0 40.0          # Solar azimuth angle
umu  -1.0 -0.5  -0.2 -0.1 # Output cosine of polar angle
phi 0.0 45. 90. 135. 180.0 225. 270.0 # Output azimuth angles
```

More examples with output are found in the ‘examples’ directory.

2.1.2 The uvspec output

The uvspec output differs for the different solvers. It may be controlled to some degree using `output_user`.

2.1.2.1 DISORT, SDISORT and SPSPDISORT

For the `disort`, `sdisort` and `spsdisor` solvers `uvspec` outputs one block of data to standard output (stdout) for each wavelength. The format of the block is

```
lambda edir edn eup uavgdir uavgdn uavgup
```

if `umu` is not specified. If `umu` is specified the format of the block is

```
lambda edir edn eup uavgdir uavgdn uavgup
umu(0) u0u(umu(0))
umu(1) u0u(umu(1))
.      .
.      .
```

If both `umu` and `phi` is specified the output format of each block is

```
lambda edir edn eup uavgdir uavgdn uavgup
                                phi(0)    ...    phi(m)
umu(0) u0u(umu(0))  uu(umu(0),phi(0)) ... uu(umu(0),phi(m))
umu(1) u0u(umu(1))  uu(umu(1),phi(0)) ... uu(umu(1),phi(m))
.      .            .
.      .            .
umu(n) u0u(umu(n))  uu(umu(n),phi(0)) ... uu(umu(n),phi(m))
```

and so on for each wavelength.

2.1.2.2 TWOSTR

The format of the output line for the `twostr` solver is

```
lambda edir edn eup uavg
```

for each wavelength.

2.1.2.3 POLRADTRAN

The output from the `polradtran` solver depends on the number of Stokes parameters, `polradtran_nstokes`. For each wavelength the output block is

```
lambda down_flux(1) up_flux(1) ... down_flux(is) up_flux(is)
```

if `phi` is not specified. Here `is` is the number of Stokes parameters specified by `polradtran_nstokes`. If `phi` and `umu` are specified the block is

```
lambda down_flux(1) up_flux(1) ... down_flux(is) up_flux(is)
                                phi(0)                        ...      phi(m)

Stokes vector I
umu(0) u0u(umu(0))   uu(umu(0),phi(0)) ... uu(umu(0),phi(m))
umu(1) u0u(umu(1))   uu(umu(1),phi(0)) ... uu(umu(1),phi(m))
.      .              .
.      .              .
umu(n) u0u(umu(n))   uu(umu(n),phi(0)) ... uu(umu(n),phi(m))
Stokes vector Q
.      .              .
.      .              .
```

Note that `polradtran` outputs the total (=direct+diffuse) downward flux. Also note that `u0u` is always zero for `polradtran`.

2.1.2.4 Description of symbols

In the above output blocks the symbols used have the following meaning.

<code>cmu</code>	Computational polar angles from <code>polradtran</code> .
<code>down_flux</code> , <code>up_flux</code>	The total (direct+diffuse) downward (<code>down_flux</code>) and upward (<code>up_flux</code>) irradiances. Same units as extraterrestrial irradiance.
<code>lambda</code>	Wavelength (nm)
<code>edir</code>	Direct beam irradiance (same units as extraterrestrial irradiance, e.g mW/(m ² nm) if using the 'atlas3' spectrum in the 'data/solar_flux' directory.)
<code>edn</code>	Diffuse down irradiance, i.e. total minus direct beam (same units as <code>edir</code>).
<code>eup</code>	Diffuse up irradiance (same units as <code>edir</code>).
<code>uavg</code>	The mean intensity. Proportional to the actinic flux: To obtain the actinic flux, multiply the mean intensity by 4 pi (same units as <code>edir</code>).

<code>uavgdir</code>	Direct beam contribution to the mean intensity. (same units as <code>edir</code>).
<code>uavgdn</code>	Diffuse downward radiation contribution to the mean intensity. (same units as <code>edir</code>).
<code>uavgup</code>	Diffuse upward radiation contribution to the mean intensity. (same units as <code>edir</code>).
<code>u0u</code>	The azimuthally averaged intensity at <code>numu</code> user specified angles <code>umu</code> . (units of e.g. $\text{mW}/(\text{m}^2 \text{ nm sr})$ if using the ‘ <code>atlas3</code> ’ spectrum in the ‘ <code>data/solar_flux</code> ’ directory.)
<code>uu</code>	The radiance (intensity) at <code>umu</code> and <code>phi</code> user specified angles. (units of e.g. $\text{mW}/(\text{m}^2 \text{ nm sr})$ if using the ‘ <code>atlas3</code> ’ spectrum in the ‘ <code>data/solar_flux</code> ’ directory.)
<code>uu_down, uu_up</code>	The downwelling and upwelling radiances (intensity) at <code>cmu</code> and <code>phi</code> angles. (units of e.g. $\text{mW}/(\text{m}^2 \text{ nm sr})$ if using the ‘ <code>atlas3</code> ’ spectrum in the ‘ <code>data/solar_flux</code> ’ directory.)

The total downward irradiance is given by

$$\text{irr_down} = \text{edir} + \text{edn}$$

The total mean intensity is given by

$$\text{uavg} = \text{uavgdir} + \text{uavgdn} + \text{uavgup}$$

If `deltam` is on it does not make sense to look at the individual contributions to `uavg` since they are delta-M scaled.

2.1.3 Complete description of input parameters

The `uvspec` input file consists of single line entries, each making up a complete input to the `uvspec` program. First on the line comes the parameter name, followed by one or more parameter values. The parameter name and the parameter values are separated by white space. Filenames are entered without any surrounding single or double quotes. Comments are introduced by a `#`. Blank lines are ignored. The order of the lines is not important, with one exception: if the same input option is used more than once, the second one will usually over-write the first one. Be aware that also option in another `included` input file will overwrite options specified before.

The various input parameters are described in detail below.

`aerosol_angstrom`

Scale the aerosol optical depth using the Ångström formula. Specify the Ångström alpha and beta coefficients. The optical thickness defined here is the integral from the user-defined `altitude` to TOA (top of atmosphere).

`aerosol_default`

Set up a default aerosol according to Shettle (1989). The default properties are a rural type aerosol in the boundary layer, background aerosol above 2km,

spring-summer conditions and a visibility of 50km. These settings may be modified with `aerosol_haze`, `aerosol_vulcan`, `aerosol_season`, and `aerosol_visibility`.

`aerosol_files`

A way to specify aerosol optical depth, single scattering albedo, and phase function moments for each layer. The file specified by `aerosol_files` has two columns where column 1 is the altitude in km. The altitudes must be the same as those specified in the `atmosphere_file`. The second column is the name of a file which defines the optical properties of the layer starting at the given altitude. The files specified in the second column must have the following format:

Column 1: The wavelength in nm. These wavelengths may be different from those in `solar_file`. Optical properties are interpolated to the requested wavelengths.

Column 2: The extinction coefficient of the layer in units km⁻¹.

Column 3: The aerosol single scattering albedo of the layer.

Column 4-(nmom+4):

The moments of the aerosol phase function.

For some simple examples see the files '`examples/AERO_*.LAYER`'. Note that if using the `rte_solver disort2` it makes good sense to make the number of moments larger than `nstr`. For `rte_solver disort` and `rte_solver polradtran` the number of moments included in the calculations will be `nstr+1`. Higher order moments will be ignored for these solvers. Please note that the uppermost line of the `aerosol_files` denotes simply the top altitude of the uppermost layer. The optical properties of this line are consequently ignored. There are two options for this line: either an optical property file with zero optical thickness is specified or "NULL" instead.

`aerosol_gg_file`

Location of aerosol asymmetry parameter file. The file must have two columns. Column 1 is the altitude in km. The altitude grid must be exactly equal to the altitude grid specified in the file `atmosphere_file`. Column 2 is the asymmetry parameter of each layer. At present, the asymmetry parameter defined here is constant with wavelength but it is planned to introduce the wavelength dependence in the near future. Comments start with `#`. Empty lines are ignored.

`aerosol_haze`

Aerosol type in the lower 2 km of the atmosphere. Integer. See E.P. Shettle, "Models of aerosols, clouds and precipitation for atmospheric propagation studies", in "Atmospheric propagation in the uv, visible, ir and mm-region and related system aspects", AGARD Conference Proceedings (454), 1989.

- 1 Rural type aerosols.
- 4 Maritime type aerosols.
- 5 Urban type aerosols.

6 Tropospheric type aerosols.

`aerosol_moments_file`

Location of aerosol moments file, a one-column file containing an arbitrary number of Legendre terms of the phase function. The phase function $p(\mu)$ is

$$p(\mu) = \sum_{m=0}^{\infty} (2m+1) \cdot k_m \cdot P_m(\mu)$$

where k_m is the m 'th moment and $P_m(\mu)$ is the m 'th Legendre polynomial. If not specified, a Henyey-Greenstein phase function is assumed where the asymmetry parameter g is either a default value depending on the aerosol type or may be specified using `aerosol_set_gg`.

`aerosol_refrac_file`

File containing the wavelength-dependent refractive index of the aerosol. Three columns are expected: wavelength [nm] and the real and imaginary parts of the refractive index. Together with `aerosol_sizedist_file` this forms the input to Mie calculations of the aerosol optical properties. Please note that only the single-scattering albedo, the scattering phase function, and the wavelength-dependence of the extinction coefficient are affected by the Mie calculation while the absolute value of the extinction coefficient is taken from other sources; generally, the extinction coefficient at the first *internal* wavelength is taken from whatever is available (either default Shettle (1989) or user-defined); the extinction at all other wavelengths is scaled according to the Mie calculation. For this reason, the absolute numbers are not relevant - only the shape of the size distribution matters. In detail: If the aerosol properties are defined using the refractive index and the size distribution, the wavelength dependence of the optical properties is determined by Mie theory. At present there are at least three ways to define the absolute value of the optical thickness: (1) `visibility` defines the profile at the first *internal* wavelength; for a monochromatic calculation and in correlated-k mode, the first *internal* wavelength equals the first wavelength output by `uvspec`; for spectral calculations, the first wavelength might be a little bit smaller than the first wavelength output by `uvspec`; (2) `aerosol_tau_file` defines the optical thickness profile at the first *internal* wavelength; or (3) absolute optical thickness and wavelength-dependence are defined by `aerosol_angstrom`. *It is recommended to avoid this option and rather to calculate the aerosol optical properties externally e.g. with mie and to pass them to uvspec with aerosol_files.*

`aerosol_refrac_index`

Wavelength-independent refractive index of the aerosol; if wavelength-dependence is required, use `aerosol_refrac_file` instead. Together with `aerosol_sizedist_file` this forms the input to Mie calculations of the aerosol optical properties. Please see the description of `aerosol_refrac_file` to learn how the optical properties are set up. *It is recommended to avoid this option and rather to calculate the aerosol optical properties externally e.g. with mie and to pass them to uvspec with aerosol_files.*

aerosol_scale_ssa

Scale the aerosol single scattering albedo for all wavelengths and altitudes with a positive number. If the resulting scaled single scattering albedo is larger than 1 it is set to 1.

aerosol_scale_tau

Scale the aerosol extinction for all wavelengths and altitudes with a positive number.

aerosol_set_gg

Set the aerosol asymmetry parameter for all wavelengths and altitudes to a constant value between -1.0 and 1.0.

aerosol_set_ssa

Set the aerosol single scattering albedo for all wavelengths and altitudes to a constant value between 0.0 and 1.0.

aerosol_set_tau

Set the aerosol optical thickness for all wavelengths and altitudes to a constant value. The optical thickness defined here is the integral from the user-defined **altitude** to TOA (top of atmosphere).

aerosol_set_tau550

Set the aerosol optical thickness at 550nm. Other wavelengths are scaled accordingly. Note that this option requires for technical reasons that the wavelength interval defined by **wavelength** does contain 550nm. The optical thickness defined here is the integral from the user-defined **altitude** to TOA (top of atmosphere).

aerosol_season

Specify season to get appropriate aerosol profile.

- 1 Spring-summer profile.
- 2 Fall-winter profile.

aerosol_sizedist_file

Aerosol size distribution. Two columns are expected: The radius [micron] and the particle number. Together with **aerosol_refrac_index** or **aerosol_refrac_file** this forms the input to Mie calculations of the aerosol optical properties. Please note that only the single-scattering albedo, the scattering phase function, and the wavelength-dependence of the extinction coefficient are affected by the Mie calculation while the absolute value of the extinction coefficient is taken from other sources; generally, the extinction coefficient at the first *internal* wavelength is taken from whatever is available (either default Shettle (1989) or user-defined); the extinction at all other wavelengths is scaled according to the Mie calculation. For this reason, the absolute numbers are not relevant - only the shape of the size distribution matters. For details see also the description of **aerosol_refrac_file**. *It is recommended to avoid this option and rather to calculate the aerosol optical properties externally e.g. with mie and to pass them to uvspec with aerosol-files.*

aerosol_ssa_file

Location of aerosol single scattering albedo file. The file must have two columns. Column 1 is the altitude in km. The altitude grid must be exactly equal to the altitude grid specified in the file **atmosphere_file**. Column 2 is the single scattering albedo of each layer. At present, the single scattering albedo defined here is constant with wavelength but it is planned to introduce the wavelength dependence in the near future. Comments start with **#**. Empty lines are ignored.

aerosol_tau_file

Location of aerosol optical depth file. The file must have two columns. Column 1 is the altitude in km. The altitude grid must be exactly equal to the altitude grid specified in the file **atmosphere_file**. Column 2 is the aerosol optical depth of each layer. To get wavelength dependence use the parameter **aerosol_angstrom**. Comments start with **#**. Empty lines are ignored.

aerosol_visibility

Visibility in km.

aerosol_vulcan

Aerosol situation above 2 km. Integer.

- 1 Background aerosols.
- 2 Moderate volcanic aerosols.
- 3 High volcanic aerosols.
- 4 Extreme volcanic aerosols.

albedo The Lambertian surface albedo, a number between 0.0 and 1.0, constant for all wavelengths. For wavelength dependent surface albedo use **albedo_file**. The default albedo is 0.0.

albedo_file

Location of surface albedo file for wavelength dependent surface albedo. The file must have two columns. Column 1 is the wavelength in nm, and column 2 the corresponding Lambertian surface albedo. The wavelength grid may be freely set. The albedo will be interpolated (linear) to the wavelength grid used for the radiation calculation. Comments start with **#**. Empty lines are ignored.

albedo_surface

With this option, the surface type and its spectral albedo is selected, see **albedo_library**.

albedo_surface surface_type

where **surface_type** is an integer starting from 1. The default value is 0, which is interpreted as black surface.

albedo_library

Albedo libraries are a collection of surface types, of which one is chosen by the option **albedo_surface** or **albedo_map**. There are two possibilities for libraries: the build in IGBP library or a user defined albedo library.

The build in library of the International Geosphere Biosphere Programme is selected with `albedo_library` IGBP. The IGBP library contains 20 surface types:

- 1 evergreen_needle_forest
- 2 evergreen_broad_forest
- 3 deciduous_needle_forest
- 4 deciduous_broad_forest
- 5 mixed_forest
- 6 closed_shrubs
- 7 open_shrubs
- 8 woody_savanna
- 9 savanna
- 10 grassland
- 11 wetland
- 12 cropland
- 13 urban
- 14 crop_mosaic
- 15 antarctic_snow
- 16 desert
- 17 ocean_water
- 18 tundra
- 19 fresh_snow
- 20 sea_ice

The first to 17th surface type are defined by the International Geosphere Biosphere Programme (IGBP), additionally there are tundra, fresh_snow, and sea_ice surface types. The spectral albedo of the ground is determined with respect to solar zenith angle, precipitable water, and cloud occurrence. The spectral resolution is like the grid of the correlated-k Fu/Liou parameterisation. This library originates from the NASA CERES/SARB Surface Properties Project, see Belward and Loveland (1996).

For creating your own albedo library use `albedo_library` path, where `path` is the path of the directory, where the albedo data is stored. The files are expected to have the names `albedo_01.dat`, `albedo_02.dat`, ... If `albedo_surface` 1 is specified the albedo from `albedo_01.dat` will be used, and so on. Each file has to have two column file: Column 1 is the wavelength in nm, and column 2 the corresponding Lambertian surface albedo. The wavelength grid may be freely set. The albedo will be interpolated (linear) to the wavelength grid used for the radiation calculation. Comments start with `#`. Empty lines are ignored. (This option is quite the same as `albedo_file`, except that it offers you an easy way to use the option `albedo_map` in combination with your albedo.files.)

`albedo_map`

Specifies a file of the surface type map, which is used in combination with `albedo_library`, `latitude`, and `longitude` in order to select the surface type relevant for the simulation. The format of the call is:

```
albedo_map file [lat_min lat_max lon_min lon_max]
```

where **file** is the location of the surface type map file. The other four optional arguments may be used, to specify the range of the map (use negative values for south and west). The default range of the map is 90 degrees south to 90 degrees north, and 0 degrees to 360 degrees east. The map is expected to be in *netCDF* format. For format discription see IGBP example and use: `ncdump libradtran/data/albedo/IGBP/SURFACE_TYPE_IGBP_10min.cdf`

For using the IGBP map, the call is `albedo_map IGBP`. This map has a resolution of 10 minutes and contains the surface types 1 to 18 defined in the `albedo_library` IGBP. Fresh snow and sea ice are not included, as their extent is too variable. (Attention: That means, that the surface type of the north pole is `ocean_water` and not `sea_ice`!)

Rounding: Locations on the pixel boundaries are interpreted as the pixel northward and eastward respectively. E.g. location 0 N, 0 E is interpreted like the pixel ranging from 0 to 10min North and from 0 to 10min East.

angstrom Still supported but obsolete. Replaced by `aerosol_angstrom`.

altitude Set the bottom level in the model atmosphere provided in `atmosphere_file` to be at altitude (km).

```
altitude 0.73 # Altitude of IFU, Garmisch-Partenkirchen
              # Be aware, for this to work the atmosphere
              # file must start at 0 km.
```

The profiles of pressure, temperature, molecular absorbers, ice and water clouds are cut at the specified altitude. The aerosol profile is not affected by **altitude** but starts right from the model surface. This is a convenient way for the user to calculate the radiation at other altitudes than sealevel. Note that **altitude** is very different from **zout** where the radiation is calculated at an altitude of **zout** over the surface. E.g. to calculate the radiation field 1 km above the surface at 0.73 above sealevel, one would specify '`altitude 0.73`' and '`zout 1.0`'.

A second optional argument may be given to **altitude** as e.g.

```
altitude 0.73 0.5
```

Here the bottom level will be at 0.73 km and the vertical resolution of the model atmosphere will be redistributed to have a spacing between levels specified by the second number, here 0.5 km. Be aware that specifying a fine vertical spacing will produce many layers thus increasing the computing time. Also the radiative transfer equation solvers implemented in Fortran 77 might need to have some array sizes increased (see '`src_f/DISORT.MXD`').

altitude_map

Specifies a file of the surface elevation map, which is used in combination with **latitude** and **longitude** in order to look up the surface elevation for the simulation. The format of the call is:

```
altitude_map file [lat_min lat_max lon_min lon_max]
```

where **file** is the location of the surface elevation map file. The other four optional arguments may be used, to specify the range of the map (use negative values for south and west). The default range of the map is 90 degrees south to

90 degrees north, and 0 degrees to 360 degrees east. The map is expected to be in *netCDF* format. For format discription see IGBP example and use: `ncdump libradtran/data/altitude/ELEVATION_GTOPO_10min.cdf`

For using the default GTOPO map, the call is `altitude_map GTOPO`. This map has a resolution of 10 minutes and the unit of the altitude is meter.

Rounding: Locations on the pixel boundaries are interpreted as the pixel northward and eastward respectively. E.g. location 0 N, 0 E is interpreted like the pixel ranging from 0 to 10min North and from 0 to 10min East.

`atmosphere_file`

Location of the atmospheric data file. Must have at least three columns containing the altitude, pressure, and temperature. Missing profiles are filled with 0 (e.g., if you did not specify the ozone profile, there will be no ozone absorption!), with exception of the air density which is calculated from pressure and temperature. Other traces gases may be set by `dens_file`. The columns are interpreted as follows:

1. `z(km)` Altitude above sea level in km
2. `p(mb)` Pressure in hPa
3. `T(K)` Temperature in K
4. `air_density(cm-3)`
 Air density in cm-3
5. `o3_density(cm-3)`
 Ozone density in cm-3
6. `o2(cm-3)`
 Oxygen density in cm-3
7. `h2o(cm-3)`
 Water vapour density in cm-3
8. `co2(cm-3)`
 CO2 density in cm-3
9. `no2(cm-3)`
 NO2 density in cm-3

The atmosphere is specified top-down, that is, the top level is the first line in the file, the bottom (surface) level the last line. All properties refer to model *level* *z*, not to model *layer*. It is important that the correct units are used, otherwise unpredictable results are guaranteed. Comments start with `#`. Empty lines are ignored. Please note that there is some redundancy: Assuming that air is an ideal gas, the air density can be calculated from pressure and temperature, $\text{dens} = p / kT$. `uvspec` will check if this relation is fulfilled and will issue a warning if it is not. libRadtran provides the six standard atmospheres by Anderson et al. (1986):

- `afglt` Tropical
- `afglms` Midlatitude Summer

<code>afglmw</code>	Midlatitude Winter
<code>afglss</code>	Subarctic Summer
<code>afglsw</code>	Subarctic Winter
<code>afglus</code>	U.S. Standard

brightness

Convert radiances / irradiances to brightness temperatures.

ch4_mixing_ratio

The mixing ratio of CH4 in ppm (default: 1.6 ppm).

cloud_fraction_file

File containing a cloud fraction profile. Two columns are expected: altitude [km] and cloud fraction, including ice and water clouds. If `cloud_fraction_file` is defined, effective cloud properties are calculated assuming random overlap of the cloud layers, as in ECHAM-4.

co2_mixing_ratio

The mixing ratio of CO2 in ppm; scale the profile so that the mixing ratio at the user-defined `altitude` assumes the specified value.

correlated_k

To calculate integrated shortwave or longwave irradiance, or to simulate satellite instrument channels, choose between the following correlated-k schemes:

Kato	Kato et al. (1999), shortwave; based on HITRAN 96. Please note that the bands above 2.5 micron are not very reliable which, however, has only little impact on integrated shortwave radiation.
Kato2	Kato et al. (1999), shortwave; optimized version (Seiji Kato, personal communication, 2003); please note that Kato2 only has 148 subbands (that is, calls to the <code>rte_solver</code>) compared to 575 for Kato which translates to an increase in computational speed by up to a factor of 4 with only little increase in uncertainty. The absorption data are based on HITRAN 2000. Please note that the bands above 2.5 micron are not very reliable which, however, has only little impact on integrated shortwave radiation.
Kato2.96	Kato et al. (1999), shortwave; optimized version (Seiji Kato, personal communication, 2003); similar to Kato2 but based on HITRAN96. Please note that the bands above 2.5 micron are not very reliable which, however, has only little impact on integrated shortwave radiation.
Fu	Fu and Liou (1992/93), shortwave and longwave; fast parameterization, developed for climate models.
AVHRR_KRATZ	Kratz (1995), AVHRR instrument channels

LOWTRAN Gas absorption parameterization from LOWTRAN; code adopted from SBDART (Ricchiazzi et al., 1998); please see the section on "Spectral resolution".

SBDART Identical to LOWTRAN.

If **correlated_k** KATO/KATO2/KATO2.96/FU/AVHRR_KRATZ is specified, the extraterrestrial flux is taken from internally defined files specific for each parameterization, not from **solar_file**. The output is the integrated irradiance for each band. To get e.g. integrated shortwave irradiance, simply add all bands of the Kato et al. (1999) or the Fu and Liou (1992/93) parameterization. The five AVHRR channels are weighted sums of the libRadtran output. Examples how to integrate the output in the AVHRR_KRATZ case are included in the **uvspec** self-test which is initiated with **make check**.

cox_and_munk_pcl

Pigment concentration for Cox and Munk ocean surface BRDF (in mg/m³); at present only available with **rte_solver** DISORT2. The number of streams (**nstr**) is automatically increased to 16 if cox_and_munk BRDF is switched on, to avoid numerical problems.

cox_and_munk_sal

Salinity for Cox and Munk ocean surface BRDF (in ppt); at present only available with **rte_solver** DISORT2. The number of streams (**nstr**) is automatically increased to 16 if cox_and_munk BRDF is switched on, to avoid numerical problems.

cox_and_munk_u10

Wind speed for Cox and Munk ocean surface BRDF (in m/s); at present only available with **rte_solver** DISORT2. The minimum allowed wind speed is 1 m/s because otherwise the strong specular reflection causes numerical problems. If a lower value is specified, the wind speed is automatically set to 1m/s. Also, the number of streams (**nstr**) is automatically increased to 16 if cox_and_munk BRDF is switched on, to avoid numerical problems.

crs_file May be used to specify cross sections of ozone (O3), nitrogendioxide (NO2), bromine oxide (BRO), OCLO, or HCHO to be used instead of one of those coming with libRadtran. No temperature dependence may be specified. Use as follows:

```
crs_file NO2 ../examples/no2_crs.dat
```

The NO2 or O3, BRO or OCLO or HCHO must be specified to identify the specie for which the cross section applies. The cross section file has two columns:

1. wavelength (nm)
2. cross section (cm²)

data_files_path

The path to the directory where all uvspec internal data files live, e.g. the files that are in the subdirectories of the 'data' directory that comes with the uvspec distribution. The default is **../data/**.

day_of_year

Integer, to correct the calculated radiation quantities for the Sun-Earth distance for the specified Julian day (1-365). If not given, the Earth-Sun distance is 1 AU (i.e. equinox distances), that is, no correction is applied to the extraterrestrial irradiance `solar_file`.

deltam Turn delta-M scaling on/off. Set to either **on** or **off**. Note that for the `rte_solver disort2` delta-M scaling is hardcoded to be always on.

dens_column

Set the total column of a density profile. The column is integrated between the user-defined altitude and TOA (top of atmosphere). The syntax is

```
dens_column specie column [unit]
```

where `specie` is one of O3, O2, H2O, CO2, NO2, BRO, OCLO, or HCHO, see also `dens_file`. The second argument is the total column value, and the optional third argument is the unit, in which the column is given. The unit can be DU (Dobson units) or CM_2 (molecules/cm²). The default units are DU for O3, and CM_2 for all other gases.

```
dens_column NO2 1.2 DU
```

Please pay attention, that the options `dens_column` and `mixing ratio` can not be used both for one gas species.

dens_file

Specify density profiles (or matrix, see below) of various trace gases to be included in the radiative transfer calculation. The entry of the input file looks like this:

```
dens_file gas_species [unit] filename
```

At the moment following gas-species are included: ozone (O3), nitrogen dioxide (NO2), water vapor (H2O), bromine oxide (BRO), OCLO, HCHO, and carbon dioxide (CO2). The gas species is identified by their abbreviations given in the parenthesis above.

Unit is an optional argument, which defines the unit the density data is given in. The profiles can be given in particles per cm³ (cm-3), in particles per m³ (m-3), as volume mixing ratio (vmr), as mass mixing ratio (mmr), or as relative humidity (only for water). The default for unit is cm-3. The model expects a density file with two columns:

1. `z(km)` Altitude above sea level in km.
2. `density of trace gas [in the specified unit]`
The density of the trace gas

The height grid may be different from the `atmosphere file`. All densities inside the range of the `dens_file` are replaced. For all other heights the values from the `atmosphere_file` are used. If the density is specified as -1 at a level, the value from `atmosphere_file` is used.

To scale the profile to a total column value use `dens_column`.

For airmass factor calculations it is for some species necessary to account for the variation of the profile with `sza`. This may be accomplished by specifying a `dens_file` in the following format:

```

0.0      SZA1      SZA2 ...
z(1)     dens(1,1) ...
z(2)     .         .
.         .         .

```

where `z(i)` are the altitude levels above sea level in km, `SZA` is the solar zenith angle in degrees, and `dens` is the density [in the specified unit] of the trace gases as function of solar zenith angle and altitude. The matrix may only be specified for one specie. It may however be combined with profiles of other species. For examples see the `examples` directory. A density matrix can only be used in connection with `rte_solver sdisort`!

`f11_mixing_ratio`

The mixing ratio of F11 in ppm (default: 0.000268 ppm).

`f12_mixing_ratio`

The mixing ratio of F12 in ppm (default: 0.000503 ppm).

`f22_mixing_ratio`

The mixing ratio of F22 in ppm (default: 0.000105 ppm).

`filter_function_file`

If specified, the calculated spectrum is multiplied with a filter function defined in '`filter_function_file`'. The file must contain two columns. Column 1 is the wavelength, in nm. Column 2 is the corresponding filter function value. Comments start with `#`. Empty lines are ignored. In combination with `output sum` or `output integrate` this option is useful e.g. to calculate weighted irradiances or actinic fluxes or to simulate broadband or satellite observations.

fisot Specifies that isotropic illumination is used at top-boundary instead of beam source. Useful for those who want to calculate the reflectance for a homogeneous or inhomogeneous atmosphere. The intensity is still set by `solar_file`.

flexstor Output is in flexstor format. May not be combined with `header`. Also, does not currently work when `umu` and/or `phi` is specified.

`h2o_mixing_ratio`

The mixing ratio of H2O in ppm. Scale the profile so that the mixing ratio at the user-defined `altitude` assumes the specified value.

`h2o_precip`

Precipitable water in kg / m2 (which is approximately 1mm). The water vapor profile is scaled accordingly. The precipitable water is integrated from the user-defined `altitude` to TOA (top of atmosphere).

header Output information on some of the input parameters. May not be combined with `flexstor`.

heating_rate

Calculation of heating rates. Output is only provided at altitudes specified by **zout**. To get heating rate profiles a number of altitudes must thus be specified. Heating rates is the change of temperature with time in units of K/day always. (In order to get the change in potential temperature multiply with $(p_0/p)^{\kappa}$.) For spectral calculations the default output is a matrix:

```

0.0      zout1      zout2 ...
lambda1  heating_rates ...
lambda2      .
.           .
.           .

```

For integrated calculations (**output sum** or **output integrate**) the default output is in two columns with column 1 being the **zout**-altitudes and column 2 the heating rates. The output of **heating_rate** can also be specified with the **output_user** option. Note that heating rates are only well-behaved up to altitudes for which the respective correlated-k options are valid. E.g. about 60 km for Fu and about 80 km for Kato, Kato2, Kato2.96, and LOWTRAN.

n2o_mixing_ratio

The mixing ratio of N₂O in ppm (default: 0.28 ppm).

ic_cloudcover

Set the fraction of the horizontal sky area which is covered by clouds. When a cloud cover is specified, the result will be calculated by the independent pixel approximation (IPA), that is, as weighted average of cloudless sky and overcast sky, where the cloud properties are taken from **ic_file**, etc. Please note that, if both **wc_cloudcover** and **ic_cloudcover** are set, both must be equal.

ic_file Location of file defining ice cloud properties. The file must contain three columns. Column 1 is the altitude in km, column 2 the ice water content in grams per cubic meter, and column 3 the effective particle radius in micron. The ice water content and effective particle radius must be specified at the same altitude grid as in **atmosphere_file**. Note that the definition of cloud altitudes in **ic_file** refers to sea level, not to altitude above ground. E.g., when altitude is set to 1.63km, and the first cloud level is defined at 3km, the cloud would start at 1.37km above ground. The optical depth of a layer is calculated using information from the upper and lower levels defining the layer. Comments start with #. Empty lines are ignored. An example of an ice cloud is given in 'examples/IC.DAT'. Per default, the cloud properties are interpreted as properties at a given altitude *level*. If **ic_layer** is defined, they are interpreted as *layer* properties (please see the section about water clouds for a realistic example how the contents of the **ic_file** are converted to optical properties).

ic_files A way to specify ice cloud optical depth, single scattering albedo, and phase function moments for each layer. The file specified by **ic_files** has two columns where column 1 is the altitude in km. The altitudes must be the same as those specified in the file **atmosphere_file**. The second column is the

name of a file which defines the optical properties of the level starting at the given altitude. The files specified in the second column must have the following format:

Column 1: The wavelength in nm. These wavelengths may be different from those in `solar_file`. Optical properties are interpolated to the requested wavelengths.

Column 2: The extinction coefficient of the layer in units km⁻¹.

Column 3: The single scattering albedo of the layer.

Column 4-(nmom+4):

The moments of the scattering phase function.

Note that if using the `rte_solver disort2` it makes good sense to make the number of moments larger than `nstr`. For `rte_solver disort` and `rte_solver polradtran` the number of moments included in the calculations will be `nstr+1`. Higher order moments will be ignored for these solvers. Please note that the uppermost line of the `ic_files` denotes simply the top altitude of the uppermost layer. The optical properties of this line are consequently ignored. There are two options for this line: either an optical property file with zero optical thickness is specified or "NULL" instead.

`ic_fu_tau`

Specify if the Fu (1996) optical properties are delta-scaled or not. With `ic_fu_tau scaled` delta-scaling is switched on, with `ic_fu_tau unscaled` it is switched off. The default is with delta-scaling, as suggested in the paper (equation 3.8 and description; equations A.2a - A.2c and description). If you define a cloud only by its microphysical properties (ice water content, effective radius), delta-scaling should certainly be switched on and you do not need to read further. If, however, you want to use the Fu (1996) parameterization in combination with one of `ic_set_tau/tau550/gg/ssa` or `ic_scale_gg/ssa` it might be reasonable to switch delta-scaling off and you should make sure that you understand the following. Citing from Fu (1996): "For nonspherical particles in cirrus clouds, it is found that a simple representation of the scattering phase function through the asymmetry factor is inadequate (Fu and Takano 1994). As demonstrated in appendix A, the fraction of scattered energy residing in the forward peak, f , needs to be removed from the scattering parameters to incorporate the strong forward peak contribution in multiple scattering." Or in other words, the sharp forward peak is truncated and added to the unscattered direct radiation. The remaining phase function (excluding the sharp forward peak) can be safely approximated by a Henyey-Greenstein function. The scaling implies a reduction of the optical thickness, the asymmetry parameter, and the single scattering albedo. This reduction can be rather severe, e.g. a factor of about 3 for the optical thickness in the visible spectral range. This implies seemingly inconsistent optical properties: For identical IWC content and effective radius, `ic_properties key/young` would give an (unscaled) optical thickness about three times higher than `ic_properties fu`. The effect on the radiation field, however, will be comparable, due the consistent scaling of

optical thickness, asymmetry parameter, and single scattering albedo. If you, however, adjust the optical thickness using e.g. `ic_set_tau`, the effect on the radiation field will be completely different because the (unscaled) optical thickness by Key (2002) has a completely different meaning as the (scaled) optical thickness by Fu (1996). In such cases it might be reasonable to switch scaling off. This is a complicated and confusing topic and it is suggested that you play around a bit with the options, read the Fu (1996) paper, and make heavy use of the `verbose` feature.

`ic_fu_reff`

If `ic_fu_reff yang` is specified, the Fu (1996/98) parameterization uses the same definition of the effective radius as the Key et al. (2002) and Yang et al. (2000) parameterization; see discussion of `ic_properties`.

`ic_habit`

Ice crystal habit for the Yang et al. (2000) and Key et al. (2002) parameterizations, see also `ic_properties key/yang`. May be one of solid-column, hollow-column, rough-aggregate, rosette-4, rosette-6, plate, droxtal, and spheroid. Please note that this parameterization is only valid for a restricted size range, depending on the habit (see table 1 in Key et al. (2002)). Also, some of the habits are only available for wavelengths below 5 micron (rosette-4) while others are only available for wavelengths larger than 3 micron (droxtal, spheroid).

`ic_ipa_files`

A two-column file, defining ice cloud property files (see `ic_file`) in the first column and the corresponding weights in the second column. The radiative transfer calculation is performed independently for each cloud column and the result is the weighted average of all independent columns. If `ic_ipa_files` and `wc_ipa_files` are both defined, both must have the same columns in the same order, otherwise `uvspec` will complain.

`ic_layer` Specify ice cloud properties for model *layers* instead of *levels* (see also `ic_file`).

`ic_properties`

Defines how ice water content and effective particle radius are translated to optical properties. Possible choices are

`fu`

Parameterization by Fu (1996), Fu et al. (1998), see `ic_file`; this is the default setting. Note that this is a parameterization which has been created to calculate fluxes but not radiances. Note also that the optical properties in the solar range provided by Fu (1996) are delta-scaled properties (that is, the forward peak of the phase function is truncated and optical thickness, asymmetry parameter, and single scattering albedo are reduced accordingly). Please see the detailed discussion in the description of `ic_fu_tau`! For wavelengths up to 4 micron Fu (1996) is used while for wavelengths larger than 4 micron Fu et al. (1998) is chosen. Please note that Fu (1996) is based on ray-tracing calculations while Fu et al. (1998) is a mixture of ray-tracing and Mie calculations (which is required for the infrared wavelengths where the geometrical assumption does

not hold). Hence, both parameterizations are not fully consistent. Rather, differences of some % are to be expected in the wavelength region where both parameterizations overlap. Also, the wavelength dependence in the solar and infrared parts is treated differently: In the solar part (Fu, 1996) the optical properties are defined for wavelength bands - hence they are assumed constant within each band. In the infrared (Fu et al. 1998) they are defined at certain wavelengths and linearly interpolated in between. If you use this option, please see also the discussion of `ic_fu_tau` and `ic_fu_reff`. The allowed range for the effective radius is from 9.315 - 65.120 micro meter.

- | | |
|---------------------|---|
| <code>echam4</code> | Use the very simple two-band parameterization of the ECHAM4 climate model, described in Roeckner et al. (1996); this is probably only meaningful if you want to compare your results with ECHAM4, the two bands are 0.2 - 0.68 micron and 0.68 - 4.0 micron; within these bands, the optical properties are assumed constant. |
| <code>key</code> | Parameterization by Key et al. (2002). This parameterization can also be used to calculate radiances because it uses a double-Henyey-Greenstein phase function which better represents both forward and backward peaks. This parameterization covers the region from 0.2 to 5.0 micron and is available for the following habits: solid-column, hollow-column, aggregate, rosette-4, rosette-6, and plate. |
| <code>yang</code> | Parameterization similar to Key et al. (2002) but based on more recent single scattering calculations. Below 3.4 micron it actually equals the Key et al. (2002) parameterization while from 3.4 - 100 micron new coefficients have been calculated with much higher wavelength resolution and better accuracy. Hence, yang should give a reasonably consistent approximation from 0.2 - 5 micron, suitable for spectrally resolved calculations of radiance and irradiance. The covered range for the effective radius depends on the <code>ic_habit</code> . (In micro meter: solid columns [5.96, 84.22], hollow columns [4.97, 70.24], rough aggregate [3.55, 108.10], rosettes-4 [2.77, 45.30], rosettes-6 [2.85, 46.01], plates [4.87, 48.18], dendrites [0.45, 1.88], droxtals [9.48, 293.32], spheroids [6.58, 203.39]). Please pay attention to the Yang's definition of the effective radius of the ice habit, see below. |
| <code>baum</code> | Use ice cloud parameterization from Brian Baum's webpage, http://www.ssec.wisc.edu/~baum/Cirrus/IceCloudModels.html . The phase function is described by 100 Legendre coefficients, calculated with delta-fit from Hu et al. (2000). This parameterization covers the region from 0.4 to 2.2 micron. If very high accuracy is needed in the vicinity of the halo, the forward peak, or the backscatter peak, <code>ic_properties baum_detailed</code> is recommended. |

- baum_detailed** See `ic_properties baum`; the phase function is described with 10,000 Legendre coefficients which guarantees a more accurate representation of the Halo etc.
- mie** Use pre-calculated Mie tables; useful for `correlated_k`; the tables are expected in `data_files_path/correlated_k/./`. For spectral or pseudo-spectral (`correlated_k sbdart`) calculations, a set of pre-calculated tables is also available; the wavelength grid points of these data has been carefully selected such that the extinction cross section, single scattering albedo, and the asymmetry parameter are accurate to 1% (compared to the fully-resolved Mie calculation) for all wavelengths between 250nm and 100 micron. For spectral or pseudo-spectral calculations `wc_properties_interpolate` has to be defined explicitly to initiate the interpolation of the optical properties to the internal wavelength grid. Please note that this option may be extremely memory-consuming because for each internal wavelength a full set of Legendre moments of the phase function is stored (up to several thousands). The Mie tables are not part of the standard distribution (because of their large size) but they are freely available from <http://www.libradtran.org>. Note that a Mie calculation assumes spherical ice particles, the scattering function of which differs systematically from non-spherical particles. Hence, `ic_properties mie` is usually not representative of natural ice clouds.
- filename** Read optical properties from specified filename; file format is as produced by the mie-tool of the `libRadtran-tool` box (see `output_user cloudprp`) or by Frank Evans' `cloudprp`; for each of the internal (computational) wavelengths, a separate file is expected (this option is subject to change).

Please note also that, in contrast to spherical particles, there is no unique definition of effective size for non-spherical particles. In particular, the above parameterizations use different definitions which, however, differ only by a constant factor. Yang et al. (2000), Key et al. (2002) use the general definition

$$r_{\text{eff}} = \frac{3 \int V(h)n(h)dh}{4 \int A(h)n(h)dh}$$

where h is the maximum dimension of an ice crystal, $n(h)$ is the number of particles with maximum dimension h in the size distribution, and V and A are the volume and mean projected area of the particles, respectively. The volume and area are based on the spherical diameter with equivalent volume and the spherical diameter with equivalent projected area as defined by Yang et al. (2000). On the other hand, Fu et al. (1996,1998) use hexagonal columns and use the following definition

$$r_{\text{eff}} = \frac{\int D^2 L n(L) dL}{2 \int (DL + \frac{\sqrt{3}}{4} D^2) n(L) dL}$$

where D is the width of the ice crystal (that is, the maximum diameter of the hexagonal area) and L is the length. The integrand in the numerator is proportional to the volume while that in the denominator is proportional to the projected area. Evaluating these formulas one finds that, for the same hexagonal particle, the effective radius would be $3\sqrt{3}/4 = 1.299$ times larger following the Yang et al. (2000), Key et al. (2002) definition than the Fu (1996,1998) definition. As an example, an effective radius of $20\mu m$ with "ic_properties fu" and $1.299 \cdot 20\mu m = 26\mu m$ with "ic_properties yang" would give comparable results for hexagonal columns. To use consistent definitions of the effective radius in both parameterizations, use `ic_fu_reff yang`.

`ic_properties_interpolate`

Interpolate ice cloud optical properties over wavelength; useful for precalculated optical property files defined with `ic_properties`. Please note that this option may be extremely memory-consuming because for each internal wavelength a full set of Legendre moments of the phase function is stored (up to several thousands).

`ic_saturate`

With this option, the relative humidity inside of ice clouds can easily be changed. This option has one necessary and one optional argument:

`ic_saturate switch [relative_humidity]`

where `switch` is `on`, `off`, or `ipa`. The second optional argument determines the relative humidity in % (with respect to water!) inside the cloud. The default value is 100.

Details for independent column simulations: Using `switch on`, the air in *all* columns will be saturated, if there is a cloud in at least one of the columns. (This option should be used for stratiform clouds.) Using `switch ipa`, clouds in one column only effect *this* column and not the other columns. (This option should be used for convective cloud fields.)

`ic_scale_gg`

Scale the ice cloud asymmetry factor for all wavelengths and altitudes with a float between 0.0 and 1.0. If you use this option in combination with the ice cloud properties by Fu (1996), please make sure that you understand the explanation of `ic_fu_tau`.

`ic_scale_ssa`

Scale the ice cloud single scattering albedo for all wavelengths and altitudes with a float between 0.0 and 1.0. If you use this option in combination with the ice cloud properties by Fu (1996), please make sure that you understand the explanation of `ic_fu_tau`.

`ic_set_gg`

Set the ice cloud asymmetry factor for all wavelengths and altitudes to a float between -1.0 and 1.0. If you use this option in combination with the ice cloud properties by Fu (1996), please make sure that you understand the explanation of `ic_fu_tau`.

ic_set_ssa

Set the ice cloud single scattering albedo for all wavelengths and altitudes to a value between 0.0 and 1.0. If you use this option in combination with the ice cloud properties by Fu (1996), please make sure that you understand the explanation of **ic_fu_tau**.

ic_set_tau

Set the total ice cloud optical depth to a constant value for all wavelengths. The optical thickness defined here is the integral from the surface at the user-defined **altitude** to TOA (top of atmosphere). This option is useful only for monochromatic calculations or in wavelength regions where the optical properties of ice clouds can be considered constant, e.g. the ultraviolet region. If you use this option in combination with the ice cloud properties by Fu (1996), please make sure that you understand the explanation of **ic_fu_tau**.

ic_set_tau550

Set the ice cloud optical thickness at 550nm. Other wavelengths are scaled accordingly. The optical thickness defined here is the integral from the surface at the user-defined **altitude** to TOA (top of atmosphere). Note that this option requires for technical reasons that the wavelength interval defined by **wavelength** does contain 550nm. If you use this option in combination with the ice cloud properties by Fu (1996), please make sure that you understand the explanation of **ic_fu_tau**.

include Include a file into the uvspec input. Works exactly like the C **#include** or the Fortran **INCLUDE** statements.

molecular_tau_file

Location of molecular absorption optical depth file. Usually, molecular absorption is calculated from trace gas concentrations provided in **atmosphere_file** (scaled with **ozone_column**, etc. Use this option only if you want to specify the optical depth directly (e.g. for a model intercomparison) or for a line-by-line calculation. If a spectral **molecular_tau_file** is specified, the wavelength grid defined there is used as internal wavelength grid for the radiative transfer calculation, if not defined otherwise with **transmittance_wl_file**. **molecular_tau_file** can be either of the following three formats:

Monochromatic:

Column 1 is the altitude in km where the altitude grid must be exactly equal to the altitude grid specified in **atmosphere_file**.
Column 2 is the absorption optical depth of each layer.

Spectral, ASCII:

The first line contains the level altitudes in decreasing order; the following lines contain the wavelength [nm] in the first column and then the absorption optical depths of each layer.

Spectral, netcdf:

netcdf is a common platform independent format; the description, a library to read and write netcdf including some tools to generate netcdf is available at

<http://www.unidata.ucar.edu/packages/netcdf/>. A `molecular_tau_file` must obey certain rules; an example is available at the libRadtran homepage, `'UVSPEC.02A.afglms.cdf'`, a line-by-line spectrum of the oxygen A-Band around 760nm, calculated for the mid-latitude summer atmosphere by Anderson et al. (1986). The advantage of netcdf compared to ASCII is that it is much faster to read, and that the file is a self-contained, including data and a description of the variables and arrays. It is therefore particularly useful for line-by-line calculations where usually many spectral data points are involved.

Comments start with `#`. Empty lines are ignored.

`no_absorption`

Switch absorption off.

`no_molecular_absorption`

Switch all (molecular, aerosol, cloud, and ice cloud) absorption off. Note that this option does not affect the single scattering albedo; e.g. with an aerosol optical depth of 1 and a single scattering albedo of 0.7, the scattering optical depth will still be 0.7, even with `no_absorption`. Note, that the thermal emission of molecules is also switched off.

`no_rayleigh`

Switch Rayleigh scattering off.

`no_scattering`

Switch scattering off.

`no2_column_du`

Obsolete, use `dens_column` instead. Set the NO2 column to a given value. The column is integrated between the user-defined `altitude` and TOA (top of atmosphere). The value must be in Dobson units. If value is negative or `no2_column_du` is not specified, the NO2 column is not scaled.

`no2_column_moleccm-2`

Obsolete, use `dens_column` instead. Set the NO2 column to a given value. The column is integrated between the user-defined `altitude` and TOA (top of atmosphere). The value must be in molecules / cm². If value is negative or `no2_column_moleccm-2` is not specified, the NO2 column is not scaled.

`nscat`

The order of scattering for the `sos` radiative transfer equation solver. Default is 20. May also be used with the `sdisort` solver. If set to 1 `sdisort` will run in single scattering mode while if set in to 2, `sdisort` runs in full multiple scattering mode.

`nrefrac`

For the `rte_solver sdisort` refraction may be included by specifying `nrefrac`. If refraction is included also set parameter `refraction_file`.

0 No refraction, default.

	1	Refraction included using fast, but harsh method.
	2	Refraction included using slow, but accurate method.
nstr	Number of streams used to solve the radiative transfer equation. Default is 6 for fluxes and 16 for radiances.	
o2_mixing_ratio	The mixing ratio of O2 in ppm; scale the profile so that the mixing ratio at the user-defined altitude assumes the specified value.	
o3_crs	Choose between the following ozone cross sections.	
	Bass_and_Paur	Bass and Paur ozone cross section.
	Molina	Molina and Molina (1986) ozone cross section.
	Daumont	Ozone cross section by Daumont et al. (1992), Malicet et al. (1995). Molina and Molina is default.
ozone_column	Obsolete, use dens_column instead. Set the ozone column to a given value. The column is integrated between the user-defined altitude and TOA (top of atmosphere). The value must be in Dobson units. If value is negative or ozone_column is not specified, the ozone column is not scaled.	
output	Output processing. Choose between the following options:	
	sum	Sum output over wavelength. Useful in combination with the correlated_k option (Kato, Kato2, Kato2.96, Fu, AVHRR_KRATZ).
	integrate	Integrate output over wavelength. Useful for spectral calculations and correlated_k LOWTRAN.
	per_nm	Output is given in W/(m2 nm) or mW/(m2 nm) (W or mW is determined by the extraterrestrial spectrum.)
	per_cm-1	Output is given in W/(m2 cm-1) or mW/(m2 cm-1).
	per_ck_band	Output is given in W/m2 or mW/m2 per correlated-k band. (This option can not be used for spectral calculations and correlated_k LOWTRAN in the solar range.)
	none	No processing - output spectral information (default).
output_user	User defined output.(This option is case sensitive.) Here the user may specify the columns desired for output. Default output is the wavelength lambda , edir , edn , eup , uavgdir , uavgdn , uavgup for disort , sdisort , and spdisort , whereas the default for twostr is lambda , edir , edn , eup , uavg . The lines containing radiances and the output of rte_solver polradtran are not affected.	

lambda Wavelength in nm.
wavenumber Wave number in cm-1.
zout Output altitude in km.
edir, eglo, edn, eup, enet, esum Direct, global, diffuse downward, and diffuse upward irradiance. Netto is global - upward, sum is global + upward.
fdir, fglo, fdn, fup, f Direct, global, diffuse downward, diffuse upward, and total actinic flux.
uavgdir, uavgglo, uavgdn, uavgup, uavg Direct, global, diffuse downward, diffuse upward, and total diffuse mean intensity (= actinic flux / 4 pi).
albedo Albedo.
heat Heating rate in K/day.

There are also some atmospheric output variables, which might be useful.

p pressure [hPa], ,
T, T_d temperature [K], dewpoint temperature [K]
theta potential temperature [K]
theta_e equivalent potential temperature [K]
n_xxx number density of the gas xxx [cm-3]
rho_xxx mass density of the gas xxx [kg/m3]
mmr_xxx mass mixing ratio of the gas xxx [kg/kg]
vmr_xxx volumn mixing ratio of the gas xxx [m3/m3]
 where xxx is one of the following gases O3, O2, H2O, CO2, NO2, BRO, OCLO, or HCHO.

phi Azimuth output angles (in degrees) in increasing order. The radiance is output at **phi** and **umu**.

phi0 Azimuth angles of the sun (0 to 360 degrees). The sun is in the south, when phi0 is 0 degree. If **phi0** varies as a function of wavelength use **sza_file**.

polradtran_aziorder Order of Fourier azimuth series: 0 is azimuthially symmetric case. Default 0.

polradtran_max_delta_tau Initial layer thickness for doubling; governs accuracy, 10E-5 should be adequate. Do not go beyond half the real precision, i.e. 10e-8 for REAL*8. Default 1.e-05.

polradtran_nstokes Number of Stokes parameters

1 for I (no polarization)

2 for I,Q

3 for I,Q,U

4 for I,Q,U,V

Default is 1.

`polradtran_quad_type`

Type of quadrature used:

G gaussian

D double gaussian,

L Lobatto

E extra-angle(s), this must be used if `polradtran` is used in combination with `umu`. Will internally use Gaussian scheme (G). See also `radtran` documentation.

Default G.

`polradtran_src_code`

Radiation sources included:

0 none

1 solar

2 thermal

3 both

Default 1.

pressure The surface pressure (at the user-defined `altitude`) in hPa. The pressure profile as well as air, O₂ and CO₂ density profiles are scaled accordingly.

prndis Specify one or more integers between 1 and 7. Print various disort input and output in disort's own format. See '`disort.doc`' for more information. **Warning:** Produces a lot of output.

quiet If specified, a number of informative messages about what options are used and progress during computations are turned off. Otherwise these messages are output to stderr. See also `verbose`.

`radiosonde`

This option allows to change the temperature and pressure profile, and optionally to specify one or more density profiles in dependence of temperature and pressure. The entry in the input file looks like this:

```
radiosonde filename [gas_species] [unit] ...
```

At the moment following `gas_species` are included: ozone (O₃), nitrogen dioxide (NO₂), water vapor (H₂O), bromine oxide (BRO), OCLO, HCHO, and carbon

dioxide (CO₂). The gas species is identified by their abbreviations given in the parenthesis above.

Unit is an optional argument, which define the unit the density data is given in. The profiles can be given in particles per cm⁻³ (CM-3), in particles per m⁻³ (M-3), as volumn mixing ratio (VMR), as mass mixing ratio (MMR), or as relative humidity (RH) (only for water). The default for unit is RH for water vapour, MMR for ozone, and CM-3 for all other gases.

The radiosonde file must have (2 + number of gases) columns:

- 1 pressure in hPa
- 2 temperature in Kelvin
- 3, 4 ... density of trace gas in the specified unit

The model will calculate a new z-grid, starting at altitude and assuming a linear temperature gradient inside the layers. The air density will be recalculated according to the ideal gas law, and the density of the well mixed gases O₂ and CO₂ will be scaled accordingly. The atmospheric data above the radiosonde data is taken from the `atmosphere_file` level by level, starting at the first pressure level above the radiosonde data. The z-grid of the `atmosphere_file` in this height region is shifted suitably. Also if the density in the radiosonde file is specified as -1 at a level, the value from the `atmosphere_file` is used.

Possible calls are

```
radiosonde ../examples/radiosonde.dat
```

just in order to change the temperature and pressure profile, or

```
radiosonde ../examples/radiosonde2.dat H2O RH O3 MMR NO2
```

where water vapour density will be given as relative humidity, ozone as mass mixing ratio, and NO₂ in cm⁻³ (default).

`rayleigh_crs`

Choose between the following Rayleigh scattering cross sections.

Bodhaine Bodhaine et al (1999) Rayleigh scattering cross section.

Nicolet Nicolet (1984) Rayleigh scattering cross section.

Bodhaine et al. is default.

`rayleigh_depol`

Rayleigh depolarization factor; the Rayleigh scattering phase function is $p(\mu) = a + b * \mu^{**2}$ where $a = 1.5*(1+depol)/(2+depol)$ and $b = 1.5*(1-depol)/(2+depol)$. By default the depolarization is calculated using the expressions from Bodhaine et al. (1999).

`rayleigh_tau_file`

Location of Rayleigh scattering optical depth file. Usually, the Rayleigh scattering cross section is calculated from the air pressure provided in `atmosphere_file` (scaled with `pressure`). Use this option only if you really want to specify the optical depth directly (e.g. for a model intercomparison). The optical thickness profile may be either monochromatic or spectral. The format is exactly the same as for `molecular_tau_file`.

reflectivity

Calculate transmission / reflectivity instead of absolute quantities. For irradiances / actinic fluxes the transmission T is defined as

$$T = \frac{E}{E_0 \cos \theta}$$

where E is the irradiance / actinic flux, E_0 is the extraterrestrial flux, and θ is the solar zenith angle. The reflectivity R is defined as

$$R = \frac{\pi \cdot L}{E_0 \cos \theta}$$

where L is the radiance, E_0 is the extraterrestrial flux, and θ is the solar zenith angle. Obviously, reflectivities do not depend on Sun-Earth distance. Please note the difference to **transmittance**.

- reverse** Option for the strong and bold. Reverses the atmospheric input to the radiative transfer solvers. That is, the atmosphere is turned on the head. Yes, that is actually useful for some purposes. If you think you need this contact the author. Otherwise, do not use.
- rh_file** File that defines a profile of relative humidity. If specified, the water vapour profile in `atmosphere_file` is over-written. If -1 is specified at a level, the value from `atmosphere_file` is used.
- rpv_file** 4 column file, containing the Rahman, Pinty, and Verstraete (RPV) BDRF parameterization, Rahman et al. (1993). Bidirectional reflectance distribution functions for a variety of surfaces are given in the paper. This option is only supported with DISORT 2.0 and MYSTIC. The columns of the input file are wavelength [nm], rho0, k, and theta. The parameters are interpolated linearly to the internal wavelength grid. To make sure that the results are reasonable, specify the RPV data on a wavelength grid similar or equal to that used internally for the radiative transfer calculation! Optionally, a fifth column with a constant scaling factor may be defined which, however, is only used by `rte_solver disort2`.
- rpv_k** Constant RPV rho0, see `rpv_file`. `rpv_k` overwrites the wavelength-dependent value defined in `rpv_file`.
- rpv_rho0** Constant RPV rho0, see `rpv_file`. `rpv_rho0` overwrites the wavelength-dependent value defined in `rpv_file`.
- rpv_theta** Constant RPV theta, see `rpv_file`. `rpv_theta` overwrites the wavelength-dependent value defined in `rpv_file`.
- rpv_sigma** Constant RPV sigma, to be used for snow (Deguenther and Meerkoetter, 2000). A wavelength dependent sigma is not yet available.

<code>rpv_t1</code>	Constant RPV t_1 , to be used for snow (Deguenther and Meerkotter, 2000). A wavelength dependent sigma is not yet available.										
<code>rpv_t2</code>	Constant RPV t_2 , to be used for snow (Deguenther and Meerkotter, 2000). A wavelength dependent sigma is not yet available.										
<code>rpv_scale</code>	Apply a constant scaling factor for the RPV BRDF. Required e.g. if the the albedo should be set to a certain value. This factor is only used by <code>rte_solver</code> <code>disort2</code> .										
<code>rte_solver</code>	Set the radiative transfer equation solver to be used. Options are <table> <tr> <td><code>disort</code></td><td>The standard plane-parallel disort algorithm by Stamnes et al. (1988), version 1.3. For documentation see '<code>src_f/DISORT.doc</code>' as well as the papers and the DISORT report at ftp://climate.gsfc.nasa.gov/pub/wiscombe/Multiple_Scatt/. To optimize for computational time and memory, please adjust the parameters in <code>src_f/DISORT.MXD</code> for your application and re-compile.</td></tr> <tr> <td><code>disort2</code></td><td>Version 2 of disort. For documentation see '<code>src_f/DISORT2.doc</code>' as well as the papers and the DISORT report at ftp://climate.gsfc.nasa.gov/pub/wiscombe/Multiple_Scatt/. <code>disort2</code> has several improvements compared to its 'ancestor' <code>disort</code> 1.3. Hence we recommend to use <code>disort2</code> rather than the older version. To optimize for computational time and memory, please adjust the parameters in <code>src_f/DISORT.MXD</code> for your application and re-compile.</td></tr> <tr> <td><code>sdisort</code></td><td>Pseudospherical disort as described by Dahlback and Stamnes (1991). Double precision version. To optimize for computational time and memory, please adjust the parameters in <code>src_f/DISORT.MXD</code> for your application and re-compile.</td></tr> <tr> <td><code>spsdisort</code></td><td>Pseudospherical disort as described by Dahlback and Stamnes (1991), single precision version. Warning: it is not recommended to use <code>spsdisort</code> for really large solar zenith angles nor for cloudy conditions. For large optical thickness it is numerically unstable and may produce wrong results. To optimize for computational time and memory, please adjust the parameters in <code>src_f/DISORT.MXD</code> for your application and re-compile.</td></tr> <tr> <td><code>polradtran</code></td><td>The plane-parallel radiative transfer solver of Evans and Stephens (1991). Includes polarization. Note that polarization effects of aerosols and clouds are currently not included.</td></tr> </table>	<code>disort</code>	The standard plane-parallel disort algorithm by Stamnes et al. (1988), version 1.3. For documentation see ' <code>src_f/DISORT.doc</code> ' as well as the papers and the DISORT report at ftp://climate.gsfc.nasa.gov/pub/wiscombe/Multiple_Scatt/ . To optimize for computational time and memory, please adjust the parameters in <code>src_f/DISORT.MXD</code> for your application and re-compile.	<code>disort2</code>	Version 2 of disort. For documentation see ' <code>src_f/DISORT2.doc</code> ' as well as the papers and the DISORT report at ftp://climate.gsfc.nasa.gov/pub/wiscombe/Multiple_Scatt/ . <code>disort2</code> has several improvements compared to its 'ancestor' <code>disort</code> 1.3. Hence we recommend to use <code>disort2</code> rather than the older version. To optimize for computational time and memory, please adjust the parameters in <code>src_f/DISORT.MXD</code> for your application and re-compile.	<code>sdisort</code>	Pseudospherical disort as described by Dahlback and Stamnes (1991). Double precision version. To optimize for computational time and memory, please adjust the parameters in <code>src_f/DISORT.MXD</code> for your application and re-compile.	<code>spsdisort</code>	Pseudospherical disort as described by Dahlback and Stamnes (1991), single precision version. Warning: it is not recommended to use <code>spsdisort</code> for really large solar zenith angles nor for cloudy conditions. For large optical thickness it is numerically unstable and may produce wrong results. To optimize for computational time and memory, please adjust the parameters in <code>src_f/DISORT.MXD</code> for your application and re-compile.	<code>polradtran</code>	The plane-parallel radiative transfer solver of Evans and Stephens (1991). Includes polarization. Note that polarization effects of aerosols and clouds are currently not included.
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<code>polradtran</code>	The plane-parallel radiative transfer solver of Evans and Stephens (1991). Includes polarization. Note that polarization effects of aerosols and clouds are currently not included.										

<code>twostr</code>	The two-stream radiative transfer solver described by Kylling et al. (1995).
<code>sos</code>	A scalar pseudospherical successive orders of scattering code. Works for solar zenith angles smaller than 90 degrees. Can calculate azimuthally averaged radiances. Set <code>nscat</code> to specify the order of scattering.
<code>montecarlo</code>	The MYSTIC code, see http://www.bmayer.de/mystic.html . Note that MYSTIC is not part of the libRadtran distribution at present. However, it has been given to some users on a collaborative basis.
<code>tzs</code>	TZS stands for "thermal, zero scattering" and is a very fast analytical solution for the special case of thermal emission in a non-scattering atmosphere. Please note that TZS does only radiance calculations at top of the atmosphere.
<code>sss</code>	SSS stands for "solar, single scattering" and is an analytical single scattering approximation which might be reasonable for a optically thin atmosphere. Please note that SSS does only radiance calculations at top of the atmosphere.
<code>null</code>	The NULL solver does not solve the radiative transfer equation. However, it sets up the optical properties, and does the post-processing; useful if you are either interested in the overhead time required by a particular model input or if you are simply interested in the optical properties, as output by <code>verbose</code> .

`slit_function_file`

If specified, the calculated spectrum is convolved with the function found in the '`slit_function_file`'. The file must contain two columns. Column 1 is the wavelength, in nm, and relative to the center wavelength. Column 2 is the corresponding slit function value. It must be unity at the maximum. The wavelength steps in the slit function file must be equidistant. Comments start with `#`. Empty lines are ignored.

`solar_file`

Location of file holding the extraterrestrial spectrum. The file must contain two columns. Column 1 is the wavelength, in nm, and column 2 the corresponding extraterrestrial flux. The user may freely use any units he/she wants on the extraterrestrial flux. The wavelength grid specified defines the wavelength resolution at which results are returned. However, the wavelength range is determined by `wavelength`. `solar_file` may be omitted for thermal radiation calculations (`source thermal`) as well as `transmittance` and `reflectivity` calculations. If omitted, the output resolution equals the internal wavelength grid which the model chooses for the radiative transfer calculation. Comments start with `#`. Empty lines are ignored. Note that `solar_file` is ignored if `correlated_k` is specified.

- source** Solar or thermal source. Set to either **solar** or **thermal**. Solar radiation is per default output in $\text{W}/(\text{m}^2 \text{ nm})$ for spectral and **correlated_k** LOWTRAN calculations. For all other **correlated_k** options the output is integrated over the wavelength band.
- Thermal radiation is per default output in $\text{W}/(\text{m}^2 \text{ cm}^{-1})$, if the bandwidth is equal to 1 cm^{-1} (default for **correlated_k** LOWTRAN calculations). Otherwise the output is the integrated flux over the wavenumber interval specified by **thermal_bandwith**, **thermal_bands_file**, or by the **correlated_k** option (Kato, Kato2, Kato2.96, Fu, AVHRR.KRATZ, or Generic).
- spline** Spline interpolate to wavelengths **lambda_0** to **lambda_1** in steps of **lambda_step**, in nm. Specified as e.g.
- ```
spline 290. 365. 0.5
```
- Here, the calculated spectrum is interpolated to wavelengths 290., 290.5, 291., ..., 364.5, 365. For interpolation to arbitrary wavelengths use **spline\_file**. The specified wavelength interval must be within the one specified by **wavelength**.
- spline\_file** Spline interpolate to arbitrary wavelengths, in nm, given as a single column in file '**spline\_file**'. The specified wavelengths must be within the range specified by **wavelength**. Comments start with **#**. Empty lines are ignored.
- surface\_temperature** Surface temperature, used for thermal infrared calculations. If not specified, the temperature of the lowest atmospheric level is used as surface temperature.
- sza** The solar zenith angle. If the solar zenith angle varies with wavelength, use **sza\_file**. The default solar zenith angle is 0.0.
- sza\_file** Location of solar zenith angle file for wavelength dependent solar zenith angle. The file must have two or three columns. Column 1 is the wavelength, in nm, and column 2 the corresponding solar zenith angle. Optionally the third column may contain the corresponding solar azimuth angle. The solar azimuth angle is only needed when calculating radiances. The wavelength grid may be freely set. The solar zenith and azimuth angle will be interpolated to the wavelength grid used for the radiation calculation. Comments start with **#**. Empty lines are ignored.
- thermal\_bands\_file** File with the center wavelengths and the wavelength band intervals to be used for calculations in the thermal range. The following three columns are expected: center (or reference) wavelength, lower wavelength limit, upper wavelength limit [nm]. **thermal\_bands\_file** defines the wavelength grid for the radiative transfer calculation. The RTE solver is called for each of the wavelengths in the first column. The atmospheric (scattering, absorption, etc) properties are also evaluated at these wavelengths. For thermal radiation calculations, the Planck function is integrated over the wavelength bands defined in the second and third columns. The result will therefore be a band-integrated irradiance which does

only make sense when the `solar_file` grid equals the `thermal_bands_file` grid.

#### `thermal_bandwidth`

Specify a constant bandwidth in cm-1 for thermal calculations. The default is 1 cm-1. This option is ignored if used together with `correlated_k` KATO/KATO2/KATO2.96/FU/AVHRR\_KRATZ.

#### `transmittance`

Calculate transmittance / reflectance instead of absolute quantities. That is, set the extraterrestrial irradiance to 1 and do not correct for Sun-Earth distance:

$$T = \frac{E}{E_0}$$

where  $E$  is the irradiance / actinic flux / radiance and  $E_0$  is the extraterrestrial flux. Please note the difference to `reflectivity`.

#### `transmittance_wl_file`

Location of single column file that sets the wavelength grid used for the internal transmittance calculations. The wavelengths must be in nm. Do not use this option unless you know what you are doing. Comments start with `#`. Empty lines are ignored.

#### `umu`

Cosine of output polar angles in increasing order, starting with negative (downward) values (if any) and on through positive (upward) values. Must not have any zero values. The azimuthally averaged intensity `u0u` is output at `umu`.

#### `verbose`

If specified abundances of informative messages are output to `stderr`. To make use of this information, you may want to write the standard `uvspec` output to one file and the diagnostic messages to another. To do so, try `(./uvspec < uvspec.inp > uvspec.out) >& verbose.txt`. The irradiances and radiances will be written to `'uvspec.out'` while all diagnostic messages go into `'verbose.txt'`. See also `quiet`.

#### `wavelength`

Set the wavelength range by specifying first and last wavelength in nm. The default output wavelength grid is that defined in `solar_file`, unless `spline` is specified. Note that the radiative transfer calculations are done on an internal grid which can be influenced with `transmittance_wl_file` or `molecular_tau_file`

#### `wavelength_index`

Set the wavelengths to be selected. To be used together with predefined wavelength grids, such as `transmittance_wl_file` `molecular_tau_file` and particularly useful in combination with the `correlated_k` option where often only a specified number of wavelength bands is required. E.g., in combination with `correlated_k` AVHRR\_KRATZ, `wavelength_index 15 15` will select wavelength index 15 which corresponds to channel 4, or `wavelength_index 10 14` will select those bands required for channel 3. Indices start from 1.

**wc\_cloudcover**

Set the fraction of the horizontal sky area which is covered by clouds. When a cloud cover is specified, the result will be calculated by the independent pixel approximation (IPA), that is, as weighted average of cloudless sky and overcast sky, where the cloud properties are taken from **wc\_file**, etc. Please note that, if both **wc\_cloudcover** and **ic\_cloudcover** are set, both must be equal.

**wc\_file**

Location of file defining water cloud properties. The file must contain three columns: Column 1 is the altitude in km, column 2 the liquid water content (LWC) in grams per cubic meter, and column 3 the effective droplet radius in micron. Empty lines are ignored. Comments start with **#**. Note that the definition of cloud altitudes in **wc\_file** refers to sea level, not to altitude above ground. E.g., when altitude is set to 1.63km, and the first cloud level is defined at 3km, the cloud would start at 1.37km above ground. An example of a cloud is given in 'examples/WC.DAT'.

Per default (**wc\_layer** is not defined) the cloud properties are interpreted as properties at a given altitude *level*. The optical depth of a layer is calculated using information from the upper and lower levels defining the layer, if the LWC in both layers is not 0. If the LWC in one of them is 0, there is no cloud in this layer.

If **wc\_layer** is defined, the informations given in **wc\_file** are interpreted as *layer* properties (please see the section about water clouds for a realistic example how the contents of the **wc\_file** are converted to optical properties).

**wc\_files**

A way to specify cloud extinction coefficient, single scattering albedo, and scattering phase function for each layer. The file specified by **wc\_files** has two columns where column 1 is the altitude in km. The second column is the name of a file which defines the optical properties of the layer starting at the given altitude. The files specified in the second column must have the following format:

**Column 1:** The wavelength in nm. These wavelengths may be different from those in **solar\_file**. Optical properties are interpolated to the requested wavelengths.

**Column 2:** The extinction coefficient of the layer in units km<sup>-1</sup>.

**Column 3:** The single scattering albedo of the layer.

**Column 4-(nmom+4):**

The moments of the scattering phase function.

Note that if using the **rte\_solver disort2** it makes good sense to make the number of moments larger than **nstr**. For **rte\_solver disort** and **rte\_solver polradtran** the number of moments included in the calculations will be **nstr+1**. Higher order moments will be ignored for these solvers. Please note that the uppermost line of **wc\_files** denotes simply the top altitude of the uppermost layer. The optical properties of this line are consequently ignored. There are two options for this line: either an optical property file with zero optical thickness is specified or "NULL" instead.

**wc\_ipa\_files**

A two-column file, defining water cloud property files (see **wc\_file**) in the first column and the corresponding weights in the second column. The radiative transfer calculation is performed independently for each cloud column and the result is the weighted average of all independent columns. If **ic\_ipa\_files** and **wc\_ipa\_files** are both defined, both must have the same columns in the same order, otherwise **uvspec** will complain.

**wc\_layer** Specify cloud properties for layers instead of levels (see also **wc\_file**). If **wc\_layer** is specified, cloud properties are assumed to be constant over the layer. The layer reaches from the level, where the properties are defined in the **wc\_file** to the level above that one. For example, the following lines

| # | z     | LWC                 | R_eff |
|---|-------|---------------------|-------|
| # | (km)  | (g/m <sup>3</sup> ) | (um)  |
|   | 4.000 | 0.0                 | 0.0   |
|   | 3.000 | 1.0                 | 10.0  |

define a cloud in the layer between 3 and 4 km with sharp boundaries.

**wc\_properties**

Define how liquid water content and effective droplet radius are translated to optical properties. Possible choices are

- hu** Parameterization by Hu and Stamnes (1993); this is the default setting. Note that the parameterization is somewhat different for 'correlated\_k FU' than for all other cases because in the latter case the parameterization from the newer (March 2000) Fu and Liou code is used while otherwise the data are taken from the original Hu and Stamnes paper. Note that this parameterization has been developed to calculate irradiances, hence it is less suitable for radiances. This is due to the use of the Henyey-Greenstein phase function as an approximation of the real Mie phase function.
- echam4** Use the very simple two-band parameterization of the ECHAM4 climate model, described in Roeckner et al. (1996); this is probably only meaningful if you want to compare your results with ECHAM4, the two bands are 0.2 - 0.68 micron and 0.68 - 4.0 micron; within these bands, the optical properties are assumed constant.
- mie** Use pre-calculated Mie tables; useful for **correlated\_k**; the tables are expected in **data\_files\_path/correlated\_k/./**  
For spectral or pseudo-spectral (**correlated\_k sbdart**) calculations, a set of pre-calculated tables is also available; the wavelength grid points of these data has been carefully selected such that the extinction cross section, single scattering albedo, and the asymmetry parameter are accurate to 1% (compared to the fully-resolved Mie calculation) for all wavelengths between 250nm and 100 micron. For spectral or pseudo-spectral calculations **wc\_properties\_interpolate** has to be defined explicitly to initiate the interpolation of the optical properties to the internal wavelength grid. Please

note that this option may be extremely memory-consuming because for each internal wavelength a full set of Legendre moments of the phase function is stored (up to several thousands). The Mie tables are not part of the standard distribution (because of their large size) but they are freely available from <http://www.libradtran.org>. This is the correct option to calculate radiances, to be preferred over the Henyey-Greenstein approach of Hu and Stammes (1993).

**filename** Read optical properties from specified filename; file format is as produced by the mie-tool of the `libRadtran-tool` box (see `output_user cloudprp`) or by Frank Evans' `cloudprp`; for each of the internal (computational) wavelengths, a separate file is expected. Use only if you really know what you are doing (this option is subject to change).

#### **wc\_properties\_interpolate**

Interpolate water cloud optical properties over wavelength; useful for precalculated optical property files defined with `wc_properties`. Please note that this option may be extremely memory-consuming because for each internal wavelength a full set of Legendre moments of the phase function is stored (up to several thousands).

#### **wc\_saturate**

With this option, the relative humidity inside of water clouds can easily be changed. This option has one necessary and one optional argument:

`wc_saturate switch [relative_humidity]`

where **switch** is **on**, **off**, or **ipa**. The second optional argument determines the relative humidity in % (with respect to water!) inside the cloud. The default value is 100.

Details for independent column simulations: Using **switch on**, the air in *all* columns will be saturated, if there is a cloud in at least one of the columns. (This option should be used for stratiform clouds.) Using **switch ipa**, clouds in one column only effect *this* column and not the other columns. (This option should be used for convective cloud fields.)

#### **wc\_scale\_gg**

Scale the water cloud asymmetry factor for all wavelengths and altitudes with a float between 0.0 and 1.0.

#### **wc\_scale\_ssa**

Scale the water cloud single scattering albedo for all wavelengths and altitudes with a float between 0.0 and 1.0.

#### **wc\_set\_gg**

Set the water cloud asymmetry factor for all wavelengths and altitudes to a float between -1.0 and 1.0. This option is useful only for monochromatic calculations or in wavelength regions where the optical properties of water clouds can be considered constant, e.g. the ultraviolet range.

**wc\_set\_ssa**

Set the water cloud single scattering albedo for all wavelengths and altitudes to a float between 0.0 and 1.0. This option is useful only for monochromatic calculations or in wavelength regions where the optical properties of water clouds can be considered constant, e.g. the ultraviolet range.

**wc\_set\_tau**

Set the total water cloud optical thickness to a constant value for all wavelengths. The optical thickness defined here is the integral from the surface at the user-defined **altitude** to TOA (top of atmosphere). This option is useful only for monochromatic calculations or in wavelength regions where the optical properties of water clouds can be considered constant, e.g. the ultraviolet range.

**wc\_set\_tau550**

Set the water cloud optical thickness at 550nm. The optical thickness defined here is the integral from the surface at the user-defined **altitude** to TOA (top of atmosphere). Other wavelengths are scaled accordingly. Note that this option requires for technical reasons that the wavelength interval defined by **wavelength** does contain 550nm.

**zout**

Output altitudes in km above surface altitude. One or more altitudes may be specified in increasing magnitude. Output altitudes must be within the range defined in the **atmosphere\_file**. Note that **zout** does not restructure the atmosphere model. Hence, if you specify **zout 0.730** and have your atmosphere model in **atmosphere\_file** go all the way down to sea level, i.e. 0.0km., output is presented at 0.730km and calculations performed with an atmosphere between 0.0 and 0.730 km (and above of course). If you want calculations done for e.g. an elevated site you have to restructure the atmosphere model and make sure it stops at the appropriate altitude. This you may either do by editing the atmosphere file or by using **altitude**. Note that for **rte\_solver polradtran** the atmosphere file must contain the altitudes specified by **zout**.

**zout\_sea** like **zout**, but above sea surface

**z\_interpolate**

The profile in the **atmosphere\_file** provides the constituents of the atmosphere at the given levels. Where additional levels are introduced and in order to calculate layer properties, an assumption about the variation of the property within the layer is required. These interpolation methods can be changed by the **z\_interpolate** option. Two arguments are required, the property, and the interpolation method:

**z\_interpolate property interpolation\_method**

Properties which may be specified are:

O3, O2, H2O, CO2, NO2, BRO, OCL0, HCHO

T temperature (here linmix is not suitable)

Possible interpolation methods are:



|               |                                                                                                                                                          |
|---------------|----------------------------------------------------------------------------------------------------------------------------------------------------------|
| <b>linear</b> | The specified property varies linearly with height.                                                                                                      |
| <b>log</b>    | The specified property varies logarithmically with height. This is a reasonable option for all well mixed trace gases.                                   |
| <b>linmix</b> | This option is only possible for gas profiles. The mixing ratio of the gas (assuming a logarithmically varying air density) varies linearly with height. |

For all gas densities the default interpolation method is **linmix**, for temperature it is **linear**.

### **zout\_interpolate**

The z-grid of optical properties is determined by the **atmosphere\_file**, and, if specified, by other profile files like **dens\_file**, **rh\_file**, or **refractive\_index\_file**. Additional levels might be introduced by the **zout** option and the second argument of the **altitude** option. By default (if **zout\_interpolate** is not specified) levels introduced by the **zout** and **altitude** options will not affect the optical property profiles, that is, the optical properties are constant within the layers specified by the **atmosphere\_file** and profile files. If **zout\_interpolate** is specified, the atmospheric profiles (tracegases, temperature ...) are interpolated to the levels introduced by **zout** and **altitude**, and optical properties are determined from the interpolated atmospheric properties. If **heating\_rate** or **rte\_solver polradtran** is specified, **zout\_interpolate** will also be automatically activated. **zout\_interpolate** generally causes smoother variation of the optical properties.

## **2.2 mie**

**mie** performs Mie scattering calculations for a specified wavelength interval. It reads input from standard input, and outputs to standard output. It is normally invoked in the following way:

```
mie < input_file > output_file
```

The format of the input and output files are described below. Several realistic examples of input files are subsequently given.

**Warning:** Please note the error checking on input variables is very scarce at the moment. Hence, if you provide erroneous input, the outcome is unpredictable.

### **2.2.1 The mie input file**

The **mie** input file consists of single line entries, each making up a complete input to the **mie** program. First on the line comes the parameter name, followed by one or more parameter values. The parameter name and the parameter values are separated by white space.

Filenames are entered without any surrounding single or double quotes.

Comments are introduced by a **#**. Blank lines are ignored.

The various input parameters are described in detail below.



**distribution**

If specified the effective radius is converted into a size distribution of droplets.

`distribution distribution_type distribution_parameter`

where distribution type is one of the two following:

**GAMMA**      The Gamma distribution of cloud droplet sizes is

$$n(r) = a r^{\alpha} \exp(-b r),$$

where alpha is the distribution parameter given as second argument. (a and b are determined automatically.) The effective radius of the distribution is  $r_{\text{eff}} = (\alpha+3)/b$ . A typical value for water clouds is  $\alpha=7$ . For ice clouds a typical value is  $\alpha=1$ . A large value of alpha gives close to a monodisperse distribution.

**LOGNORMAL**

The lognormal distribution of cloud droplet sizes is

$$n(r) = a/r \exp(-[\ln(r/r_0)]^2 / [2 \sigma^2]),$$

where  $r_0$  is the logarithmic mode of the distribution (calculated automatically) and sigma is the standard deviation, which is given by second argument. The effective radius of the distribution is  $r_{\text{eff}} = r_0 \exp(2.5 \sigma^2)$ . A common value for water clouds is  $\sigma=0.35$ .

**mass\_density**

Specifies the mass density of the medium. Useful in combination with **refrac** **user** and **output\_user** **cloudprop**, as the format of **cloudprop** specifies the extinction coefficient per mass and not per volume as usual in this mie program.

**mie\_program**

Specify which Mie program to use:

**BH**      The Mie scattering program by Bohren and Hoffmann,  
<ftp://astro.princeton.edu/draine/scat/bhmie/>

**MIEVO**      The Mie scattering program by W. Wiscombe. For documentation  
 see <src.f/MIEV.doc> and the NCAR Mie report at  
[ftp://climate.gsfc.nasa.gov/pub/wiscombe/Single\\_Scatt/](ftp://climate.gsfc.nasa.gov/pub/wiscombe/Single_Scatt/).

**mimcut**      (positive) value below which imaginary refractive index is regarded as zero  
 (computation proceeds faster for zero imaginary index). Only used by **mie\_**  
**program** **MIEVO**.

**nmom**      Number of moments of the phase function to be calculated (default: 0). Only  
 possible with **mie\_program** **MIEVO**.

**output\_user**

The mie output is one line of output quantities to standard output (stdout) for each wavelength and each particle radius. With this option the user may specify the columns desired for output:

output\_user output\_1 output\_2 ... output\_n

where output\_i is one of following arguments:

- lambda**      Wavelength in nm.
- wavenumber**  
Wave number in cm-1.
- r\_eff**      particle radius in micro meter.
- refrac\_real**  
The real part of the refractive index.
- refrac\_img**  
The imaginary part of the refractive index.
- qext**      The extinction efficiency factor, if **r\_eff** was specified, or the extinction coefficient [km-1] per unit concentration [cm3/m3], if a **size\_distribution\_file** was specified. If the medium is liquid water, 1 cm3/m3 equals a liquid water content of 1g/m3 because the density of water is close to 1 g/cm3. For ice and other substances, the density has to be considered (0.917 g/cm3 for ice at 273K).
- qsca**      The scattering efficiency factor, if **r\_eff** was specified, or the extinction coefficient [km-1] per unit concentration [cm3/m3], if a **size\_distribution\_file** was specified.
- omega**      The single scattering albedo.
- gg**      The asymmetry parameter.
- spike**      To quote from Wiscombe's 'MIEV0.doc':  
(REAL) magnitude of the smallest denominator of either Mie coefficient (a-sub-n or b-sub-n), taken over all terms in the Mie series past N = size parameter XX. Values of SPIKE below about 0.3 signify a ripple spike, since these spikes are produced by abnormally small denominators in the Mie coefficients (normal denominators are of order unity or higher). Defaults to 1.0 when not on a spike. Does not identify all resonances (we are still working on that).  
Meaningless if a **size\_distribution\_file** was specified.
- pmom**      The nmom+1 moments (from 0 to nmom, see option **nmom**) of the phase function. The phase function  $p(\mu)$  is

$$p(\mu) = \sum_{m=0}^{\infty} (2m+1) \cdot k_m \cdot P_m(\mu)$$

where  $k_m$  is the m'th moment and  $P_m(\mu)$  is the m'th Legendre polynomial.

**cloudprp** This is a special option, which must be the only item behind **output\_user**, as **cloudprp** specifies a whole format of the output. In particular this option is useful, when a correlated-k wavelength grid is specified with **wavelength**. If specified, the output will be written in a format, which can be directly used by **libRadtran** or can be converted with **cloudprp2cdf.sh** into netCDF format, if **gawk** is installed, and then used by **libRadtran**. See **libRadtran-manual ic\_properties** and **wc\_properties** and there the items **mie** and **filename**.

The default output is:

```
lambda refrac_real refrac_img qext omega gg spike pmom
```

**r\_eff** The radius [micron] of the particle to calculate single scattering properties of. Used together with the wavelength information to calculate the Mie size parameter.

```
r_eff radius
```

The user can optionally specify a 2nd and 3rd argument to make a loop over several radii:

```
r_eff radius_min radius_max radius_step
```

First calculations is done with **radius\_min**, which will be increased by **radius\_step** until **radius\_max** is reached.

**refrac** Specify which refractive index to use. The following options are implemented:

**ice** The complex refractive index is taken from the **REFICE** function of W. Wiscombe.

**water** The complex refractive index is taken from the **REFWAT** function of W. Wiscombe.

**user <re> <im>**

A user defined refractive index. **re** and **im** are the real and imaginary parts (both positive numbers).

**file <filename>**

Read refractive index from a three-column file containing wavelength [nm], and the real and imaginary parts of the refractive index (both positive numbers). The Mie calculation is done for each wavelength defined here.

**size\_distribution\_file**

Specify a two column file, **r** [micron], **n(r)**, which describes a size distribution of droplets. The Mie calculation is repeated for each value of **r** found in the size distribution file, and the final result is a weighted average of these values. The user himself therefore has to choose a set of **r**'s suited for his specific purpose.

**temperature**

Ambient temperature, used to calculate the refractive indices of water and ice. Temperature dependence is only considered above 10 micron (water) and 167 micron (ice), respectively. Default: 300K.

**wavelength**

Sets the wavelength range, in nm. Specify first wavelength and last wavelength. The wavelength step is specified by **wavelength\_step**. Ignored if **refrac file** is specified.

**wavelength\_step**

The wavelength step, in nm. Ignored if **refrac file** is specified.

**wavelength\_index**

Set the wavelengths to be selected. This might be the normal wavelength grid defined by **wavelength** and **wavelength\_step** or a **correlated\_k** wavelength grid. E.g., in combination with **wavelength AVHRR\_KRATZ**, **wavelength\_index 15 15** will select wavelength index 15 which corresponds to channel 4, or **wavelength\_index 10 14** will select those bands required for channel 3. Indices start from 1.

## 2.2.2 The mie output

standard output (stdout) of the mie program is one line for each wavelength and each effective radius. The format of the output line is

```
lambda refrac_real refrac_img qext omega gg spike pmom
```

The keywords here are the same as in input option **output\_user**.

## 2.2.3 Examples of mie input files

An example of a complete input file is

```
mie_program MIEV0
refrac ice
mimcut 0.0000000001
r_mean 200.
wavelength 280. 5000.
wavelength_step 5.
```

## 2.3 integrate

**integrate** calculates the integral between limits **x\_min** and **x\_max** by interpolating the data points (**x[i]**, **y[i]**) with natural cubic splines or linear interpolation. **x\_min** and **x\_max** are the minimum and maximum values of the first column in the **input\_file**. The x-values in the first column must be in ascending order.

The different options to **integrate** are displayed when executing:

```
integrate -h
```

## 2.4 spline

**spline** interpolates discrete data points using natural cubic splines or linear interpolation. The x-values in the first column must be in ascending order.

The different options to **spline** are displayed when executing:

```
spline -h
```

## 2.5 conv

**conv** convolutes a spectrum with a given filter function.

The different options to **conv** are displayed when executing:

```
conv -h
```

## 2.6 addlevel

**addlevel** is a simple shell script to add a level to one of the existing standard profiles.

The different options to **addlevel** are displayed when executing:

```
addlevel -h
```

## 2.7 snowalbedo

**snowalbedo** calculate the diffuse and direct albedo of snow as formulated by Wiscombe and Warren (1980).

The different options to **snowalbedo** are displayed when executing:

```
snowalbedo -h
```

## 2.8 cldprp

**cldprp** calculates wavelength-dependent cloud properties using one of several parameterizations.

The different options to **cldprp** are displayed when executing:

```
cldprp -h
```

## 2.9 Geno3tab

The Perl script **Gen\_o3\_tab.pl** is used to generate a matrix of ozone values for solar zenith angle versus a chosen ratio of global irradiance at different wavelengths. The table is read by the C program **read\_o3\_tab** which, for a solar zenith angle and a measured ratio, returns the overhead ozone column. All available options are displayed when executing

```
perl Gen_o3_tab.pl --help
```

and

```
read_o3_tab -h
```

The following different types of tables may be generated.

### 2.9.1 Simple wavelength ratios with Gen\_o3\_tab

The simplest type of table is made of ratios of the global irradiance at two single wavelengths. This is the type of table described by Stamnes et al. (1991). This type of table is typically used to analyse measurements of the global irradiance from spectroradiometers. It is generated by the following command (\ is line continuation character)

```
perl Gen_o3_tab.pl --slitfunction slitfncfile --lower_lambda 305. \
--upper_lambda 340. --file table.dat
```

Here ‘slitfncfile’ is the name of the slit function file. It is a two column file where the first column is the wavelength (nm, in relative units) and the second column holds the slit function. The slit function must be normalized to unity at the center wavelength.

The generated table ‘table.dat’ is read by `read_o3_tab` for a measured ratio, `-r 10.0`, and solar zenith angle, `-s 30.0`, corresponding to the modelled ratio in the table

```
read_o3_tab -r 10.0 -s 30.0 table.dat
```

### 2.9.2 Bandpassed wavelength ratios with Gen\_o3\_tab

Instead of using single wavelengths it may be of advantage to use ratios of irradiances covering a certain wavelength range and weighted with a bandpass function. This approach may reduce problems due to changes in cloud cover and experimental uncertainties. This approach is also suitable to calculate ozone columns from multichannel, moderate bandwidth filter instruments, Dahlback (1996). Such tables are generated by

```
perl Gen_o3_tab.pl --slitfunction slitfncfile --lower_lambda 305.0 \
--upper_lambda 320.0 --file table.dat \
--bandpasslower bplow.dat --bandpassupper bpupp.dat
```

Here ‘bplow.dat’ and ‘bpupp.dat’ are the bandpass function of the lower and upper wavelength region respectively. The bandpass files have two columns. The first column is the wavelength in nm and relative units to `-lower_lambda` and `-upper_lambda`. If absolute units are specified as for filter instruments, use the `-absolute` option. The second column is the bandpass function.

The tables are read in the same way as the simple wavelength ratio tables.

## 2.10 Genwctab

The Perl script `Gen_wc_tab.pl` is used to generate a matrix of cloud optical depth for solar zenith angle versus a chosen global irradiance at a selected wavelength. The wavelength should be chosen such that it is not affected by ozone, e.g. 380 nm. The table is read by the C program `read_o3_tab` which, for a solar zenith angle and a measured irradiance, returns the overhead cloud optical depth. All available options are displayed when executing

```
perl Gen_wc_tab.pl --help
```

and

```
read_o3_tab -h
```

The following different types of tables may be generated.

### 2.10.1 Simple wavelength ratios with `Gen_wc_tab`

The simplest type of table is made of the global irradiance at a single wavelength. This is the type of table described by Stamnes et al. (1991). This type of table is typically used to analyse measurements of the global irradiance from spectroradiometers. It is generated by the following command (\ is line continuation character)

```
perl Gen_wc_tab.pl --slitfunction slitfncfile --lambda 380. \
--file table.dat
```

Here ‘`slitfncfile`’ is the name of the slit function file. It is a two column file where the first column is the wavelength (nm, in relative units) and the second column holds the slit function. The slit function must be normalized to unity at the center wavelength.

The generated table ‘`table.dat`’ is read by `read_o3_tab` for a measured global irradiance, `-r 10.0`, and solar zenith angle, `-s 30.0`, corresponding to the modelled ratio in the table. The table must be corrected for the Earth–Sun distance for the day of the measurement. This is achieved by specifying `-d 170`, where 170 is the day number. The table is generated for day 1.

```
read_o3_tab -r 10.0 -s 30.0 -d 170 table.dat
```

### 2.10.2 Bandpassed wavelength ratios with `Gen_wc_tab`

Instead of using a single wavelength it may be of advantage to use irradiances covering a certain wavelength range and weighted with a bandpass function. This approach may reduce problems due to changes in cloud cover and experimental uncertainties. This approach is also suitable to calculate cloud optical depth from multichannel, moderate bandwidth filter instruments, Dahlback (1996). Such tables are generated by

```
perl Gen_wc_tab.pl --slitfunction slitfncfile --lambda 380.0 \
--file table.dat --bandpass bp.dat
```

Here ‘`bp.dat`’ is the bandpass function of the wavelength region. The bandpass file have two columns. The first column is the wavelength in nm and relative units to `-lambda`. If absolute units are specified as for filter instruments, use the `-absolute` option. The second column is the bandpass function.

The tables are read in the same way as the simple wavelength irradiance tables.





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